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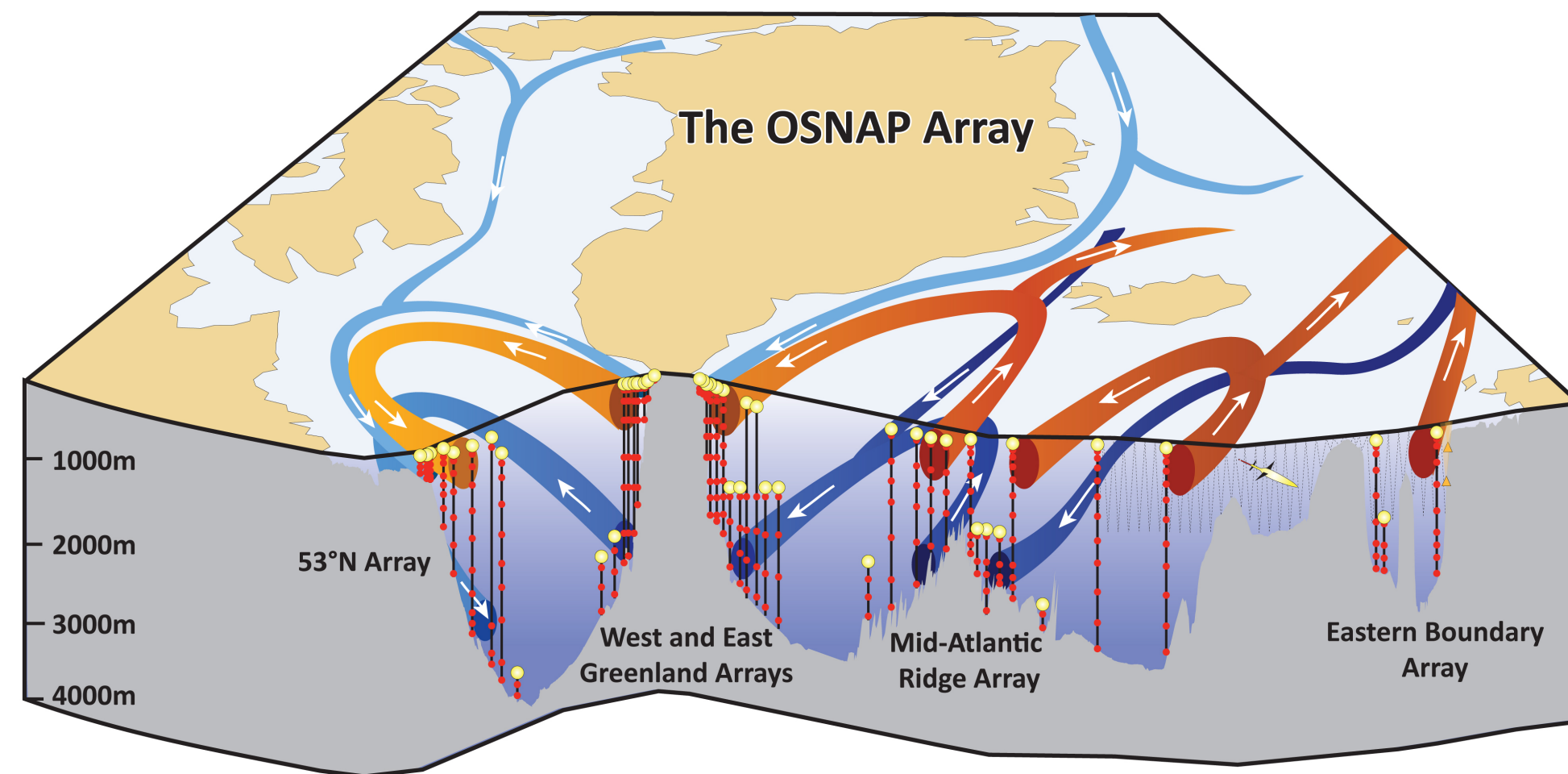
FOR COMPUTATIONAL ENGINEERING & SCIENCES

DESIGN OF AN OCEAN CLIMATE OBSERVING NETWORK IN THE SUBPOLAR NORTH ATLANTIC VIA HESSIAN UNCERTAINTY QUANTIFICATION

Nora Loose, *University of Colorado, Boulder*

Helen Pillar & Patrick Heimbach, *University of Texas at Austin*

Ocean Observing Systems



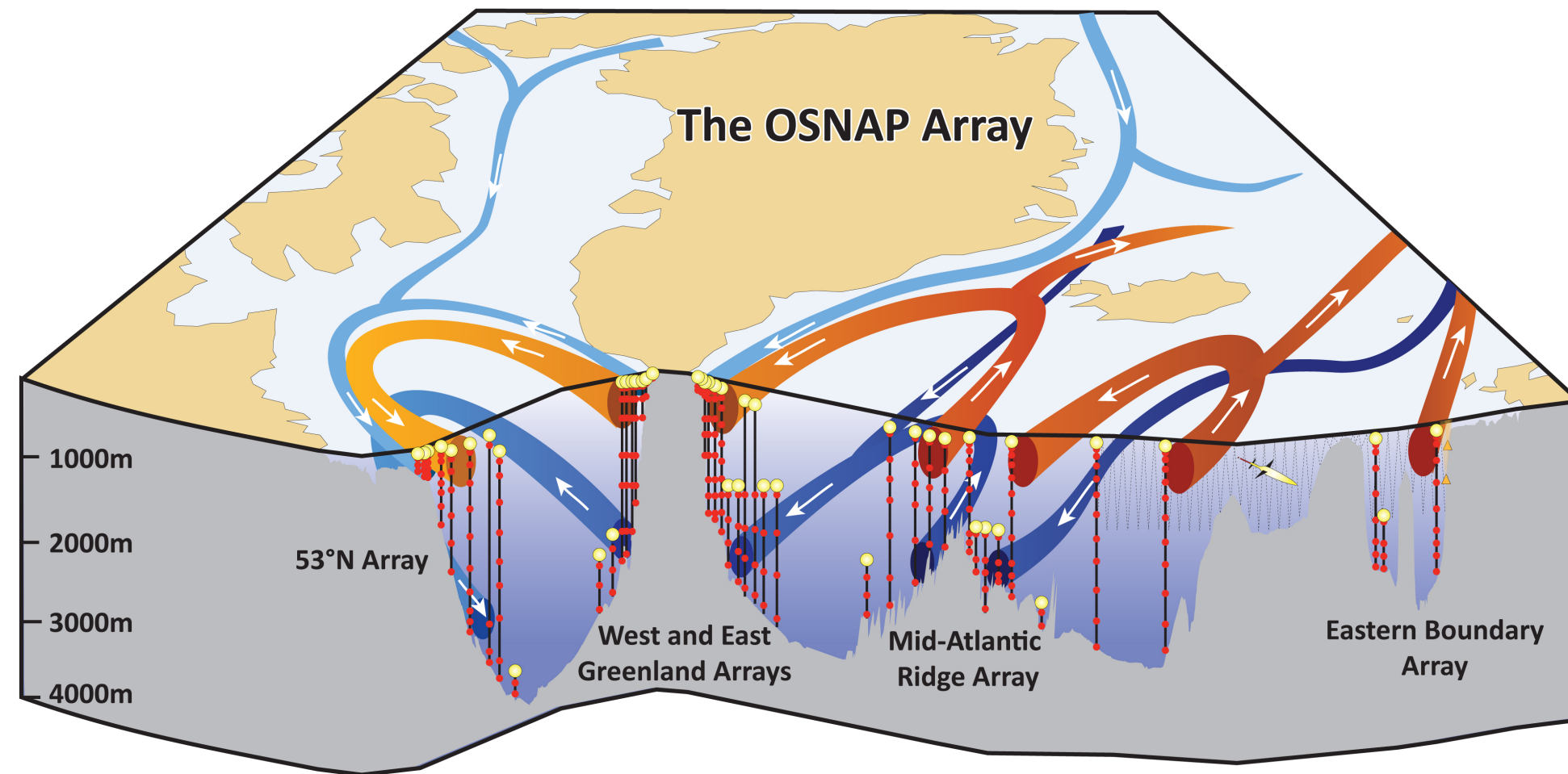
... are crucial for understanding the ocean's role in climate
... but are difficult & expensive to build and maintain

Overturning in the Subpolar North Atlantic Program (OSNAP)

<http://www.o-snap.org> (since Fall 2014)

Lozier et al., Science (2019)

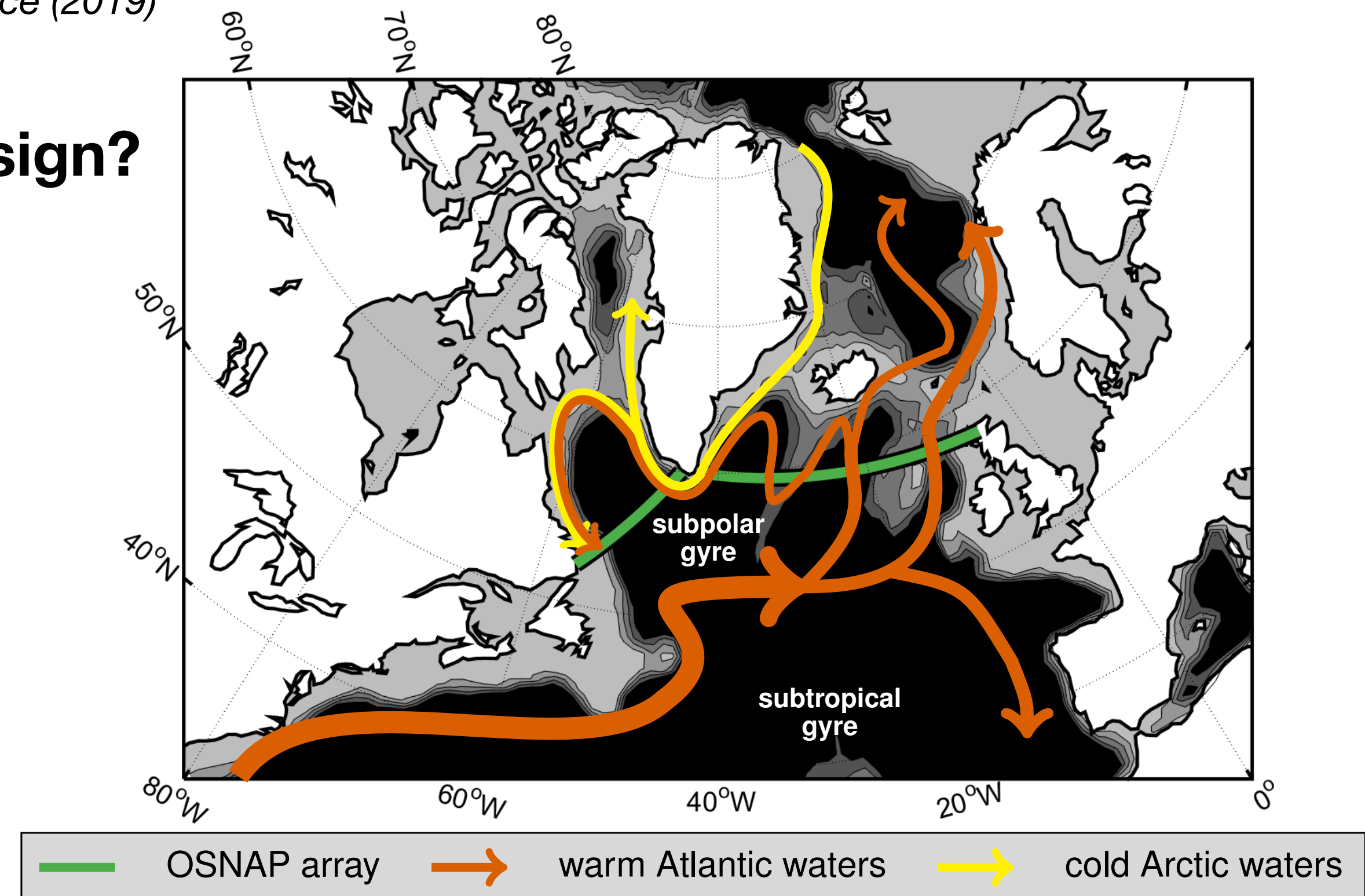
Ocean Observing Systems



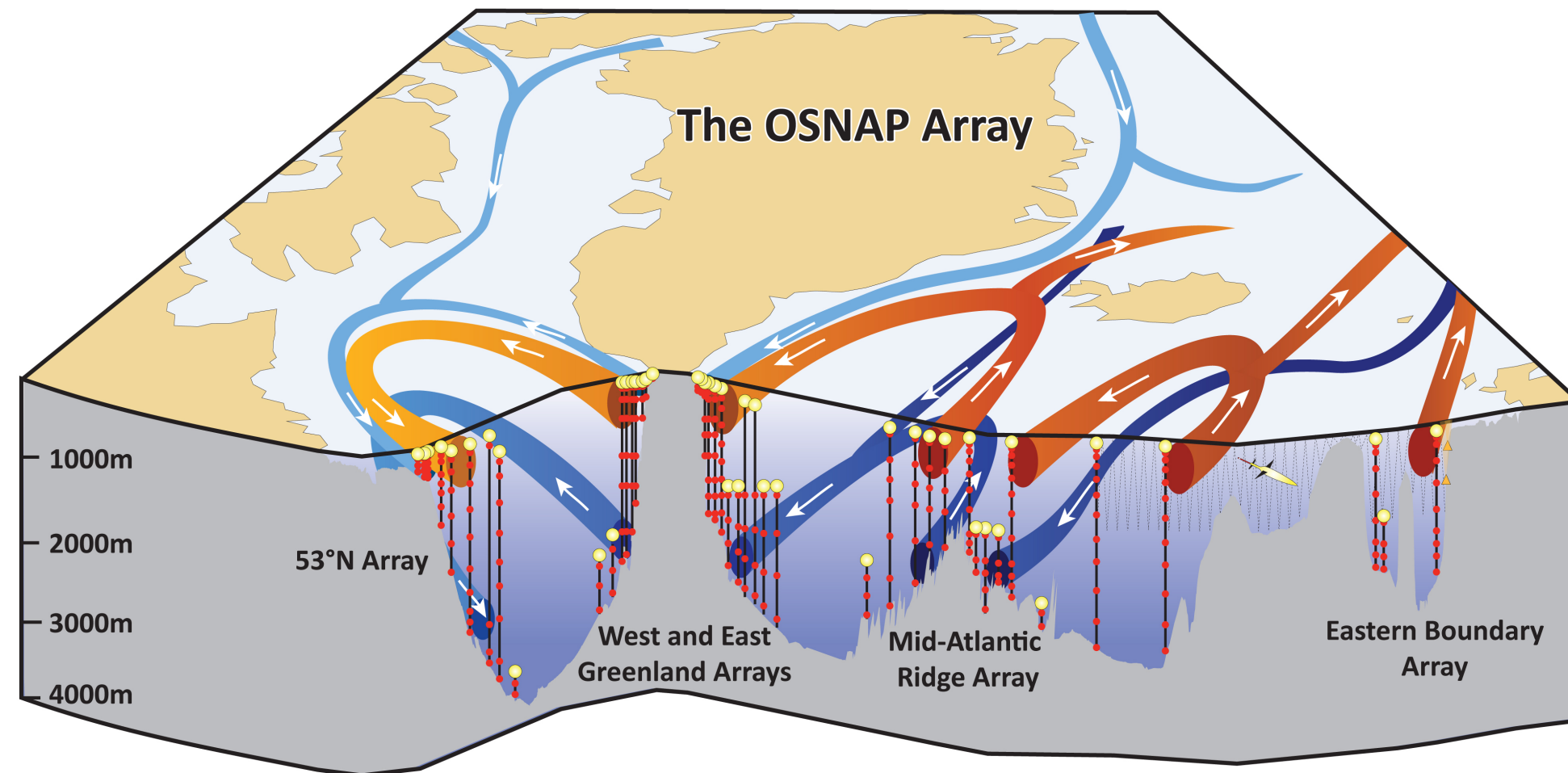
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How can ocean models inform observing system design?



Ocean Observing Systems



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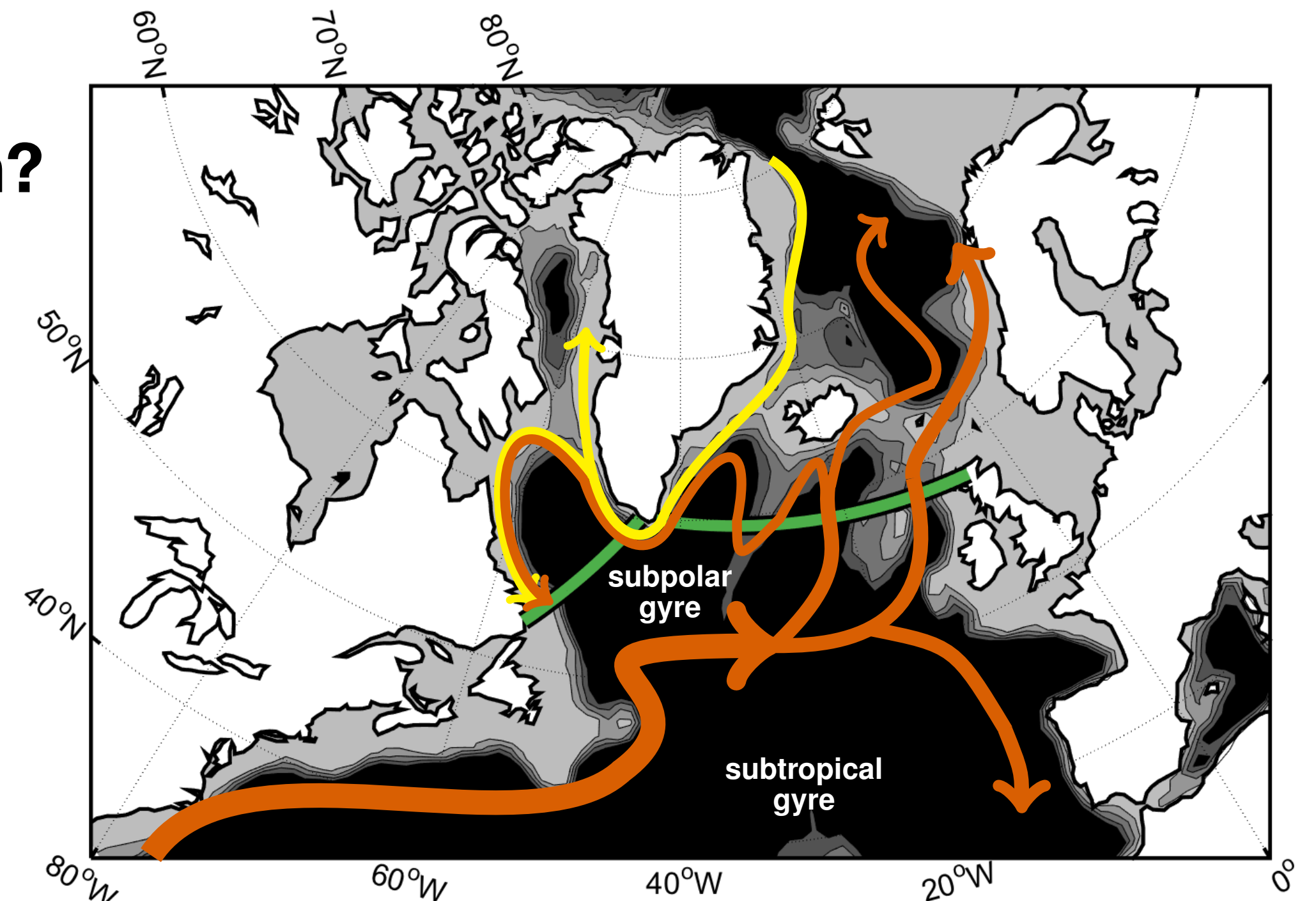
Overtuning in the Subpolar North Atlantic Program (OSNAP)

<http://www.o-snap.org>

Lozier et al., 2019

How can ocean models inform observing system design?

- Idea:**
- Use oceanic teleconnections that propagate observational constraints over long distances and can be exposed via the adjoint state
 - Combine adjoint-based estimation framework with Hessian-based uncertainty quantification



— OSNAP array → warm Atlantic waters → cold Arctic waters

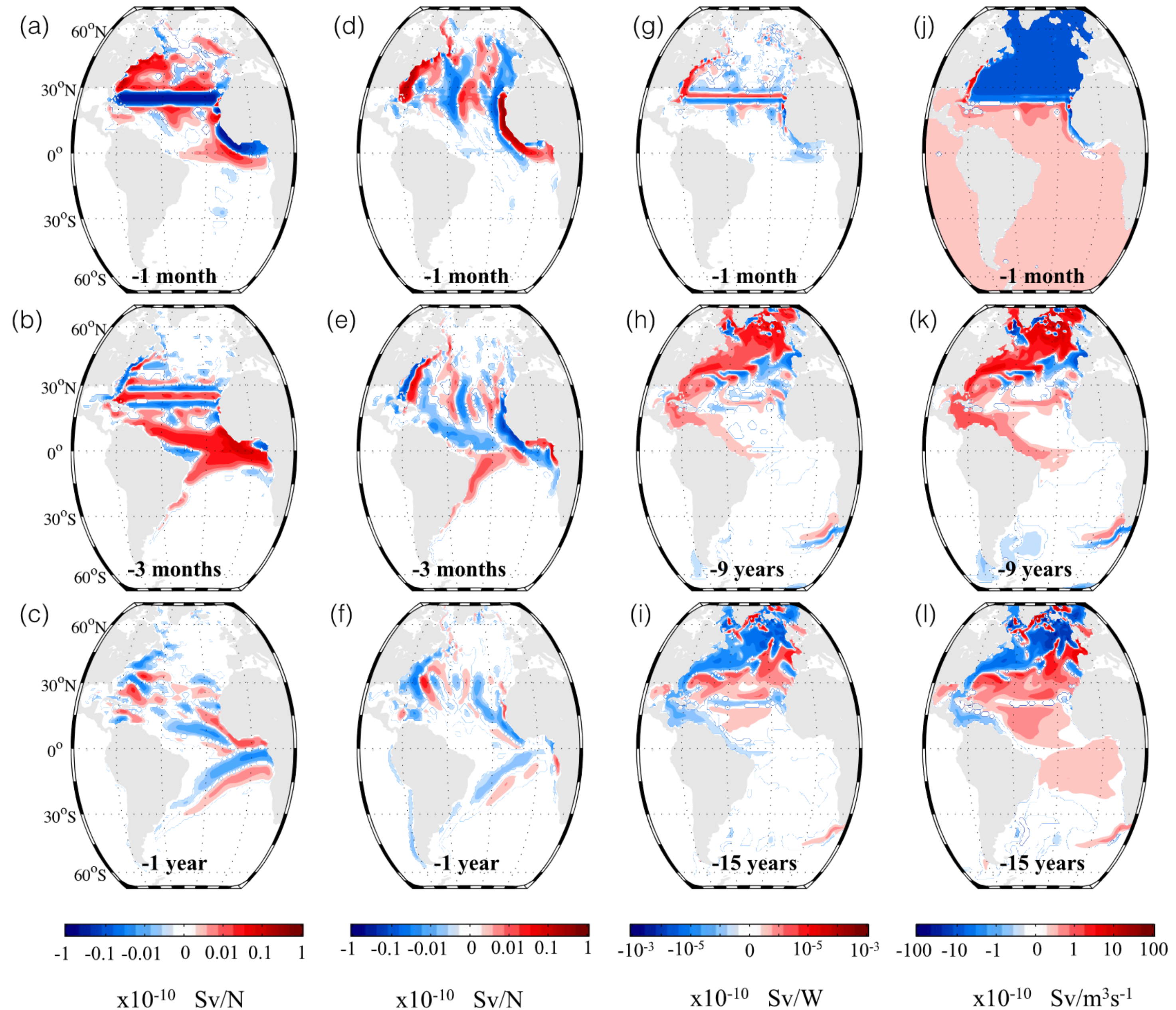
→ Quantitative design of ocean observing systems

Oceanic teleconnections

linear adjustment processes:

- oceanic Kelvin & Rossby waves
- exposed by the adjoint state as "time-reversed" waves, reflecting the sensitivity of a Quantity of Interest (QoI) to perturbations, (here: volume transport across 26N)
 - back in time, and
 - anywhere in space

Johnson & Marshall, *J. Phys. Oceanogr.* (2002)
 Heimbach et al. *Deep Sea Res.* (2011)
 Pillar et al., *J. Clim.* (2016)



Oceanic teleconnections

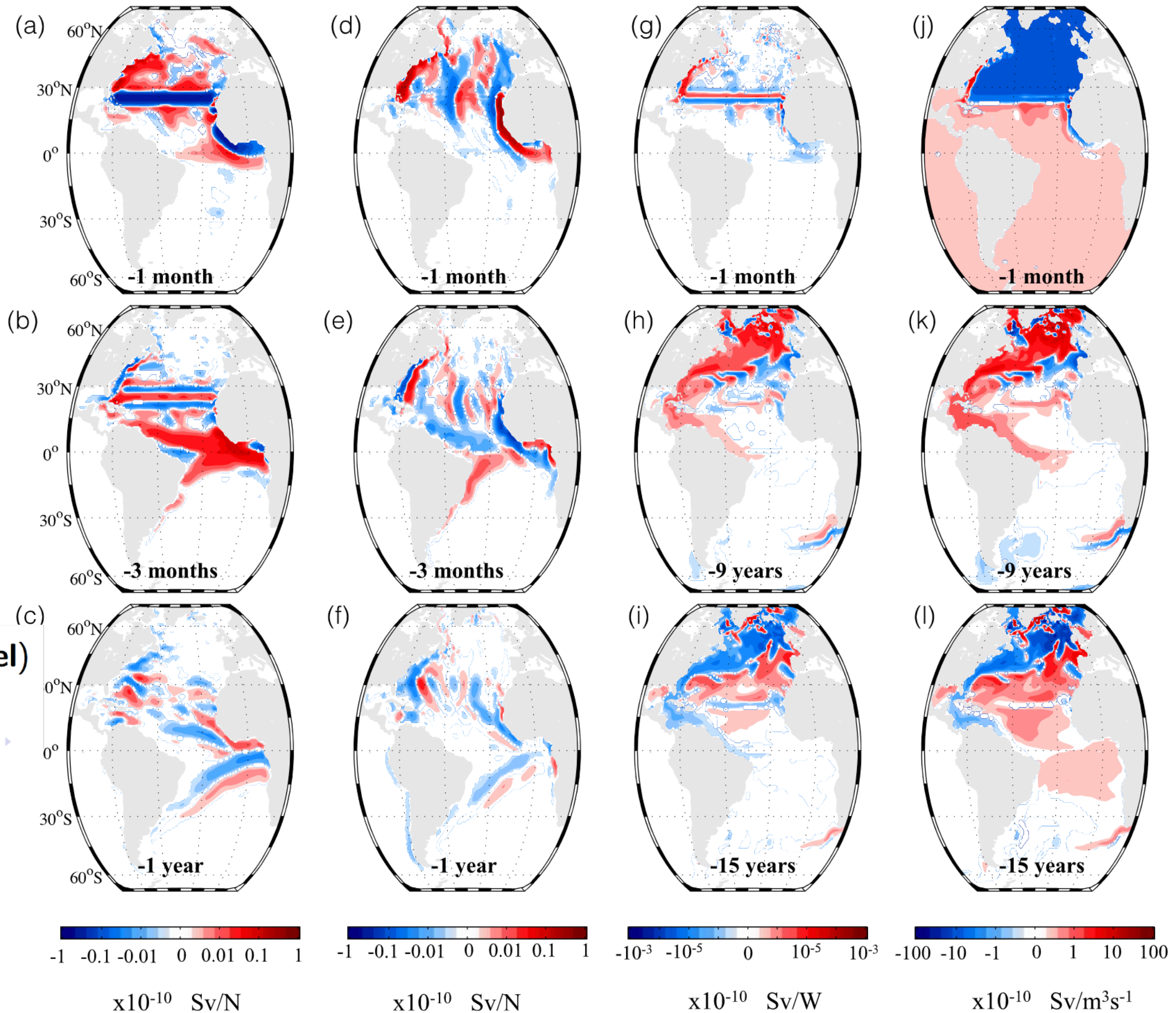
linear adjustment processes:

$$\begin{aligned} \mu_0 &= \frac{\partial J}{\partial x_0} = \sum_{1 \leq t \leq t_f} \frac{\partial x_t}{\partial x_0} \left(\frac{\partial J}{\partial x_t} \right) \\ &= \frac{\partial x_1}{\partial x_0} \left(\frac{\partial J}{\partial x_1} \right) + \frac{\partial x_1}{\partial x_0} \frac{\partial x_2}{\partial x_1} \left(\frac{\partial J}{\partial x_2} \right) \\ &\quad + \dots + \frac{\partial x_1}{\partial x_0} \dots \frac{\partial x_{t_f}}{\partial x_{t_f-1}} \left(\frac{\partial J}{\partial x_{t_f}} \right) \\ &= \mathbf{L}^T \frac{\partial J}{\partial x_1} + \mathbf{L}^T \mathbf{L}^T \frac{\partial J}{\partial x_2} + \dots + \mathbf{L}^T \dots \mathbf{L}^T \frac{\partial J}{\partial x_{t_f}} \end{aligned}$$

\mathbf{L}^T : is the adjoint model (and \mathbf{L} is the tangent linear model)

$\mu_k = \left(\frac{\partial J}{\partial x_k} \right)$: Lagrange multipliers or gradients

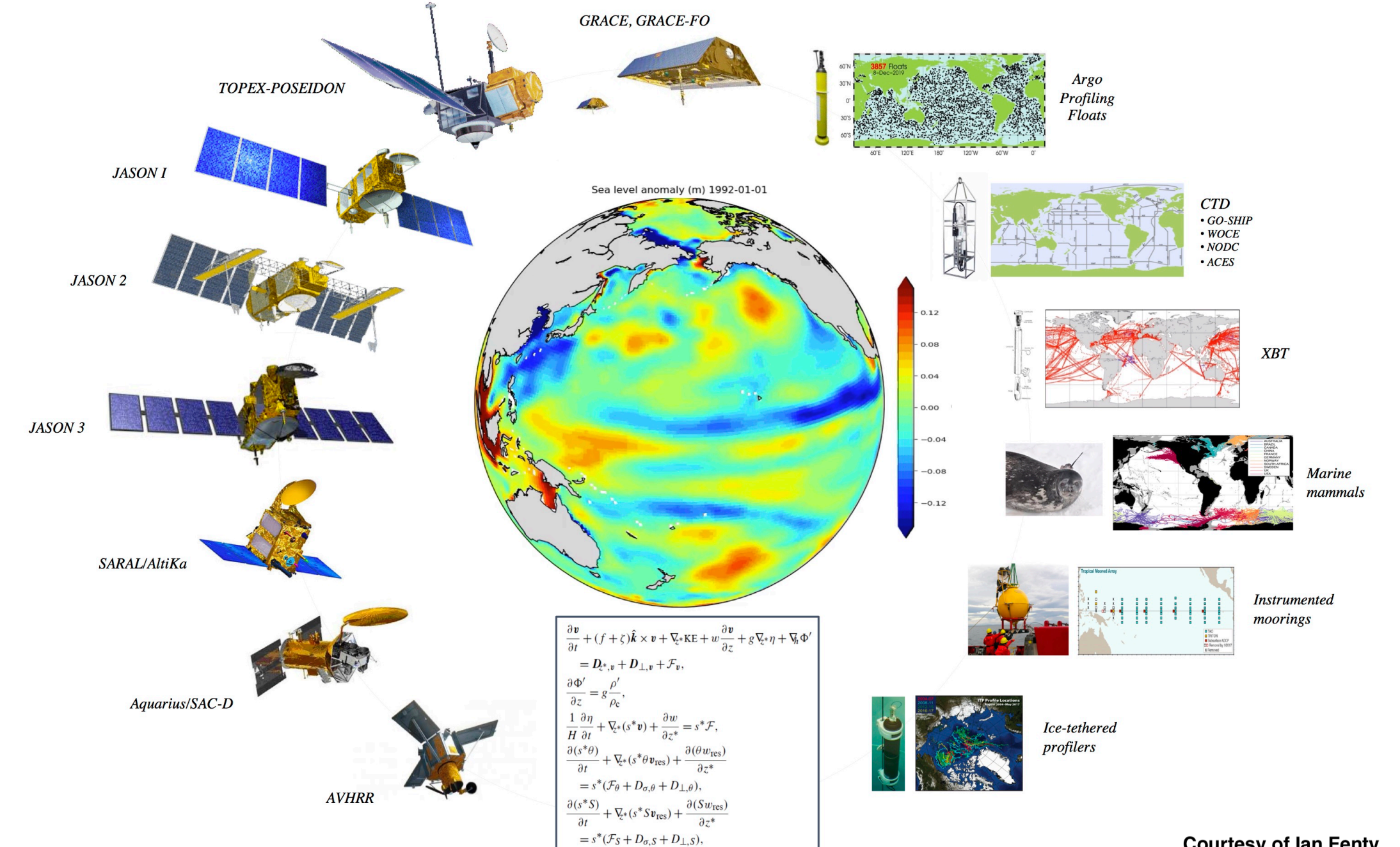
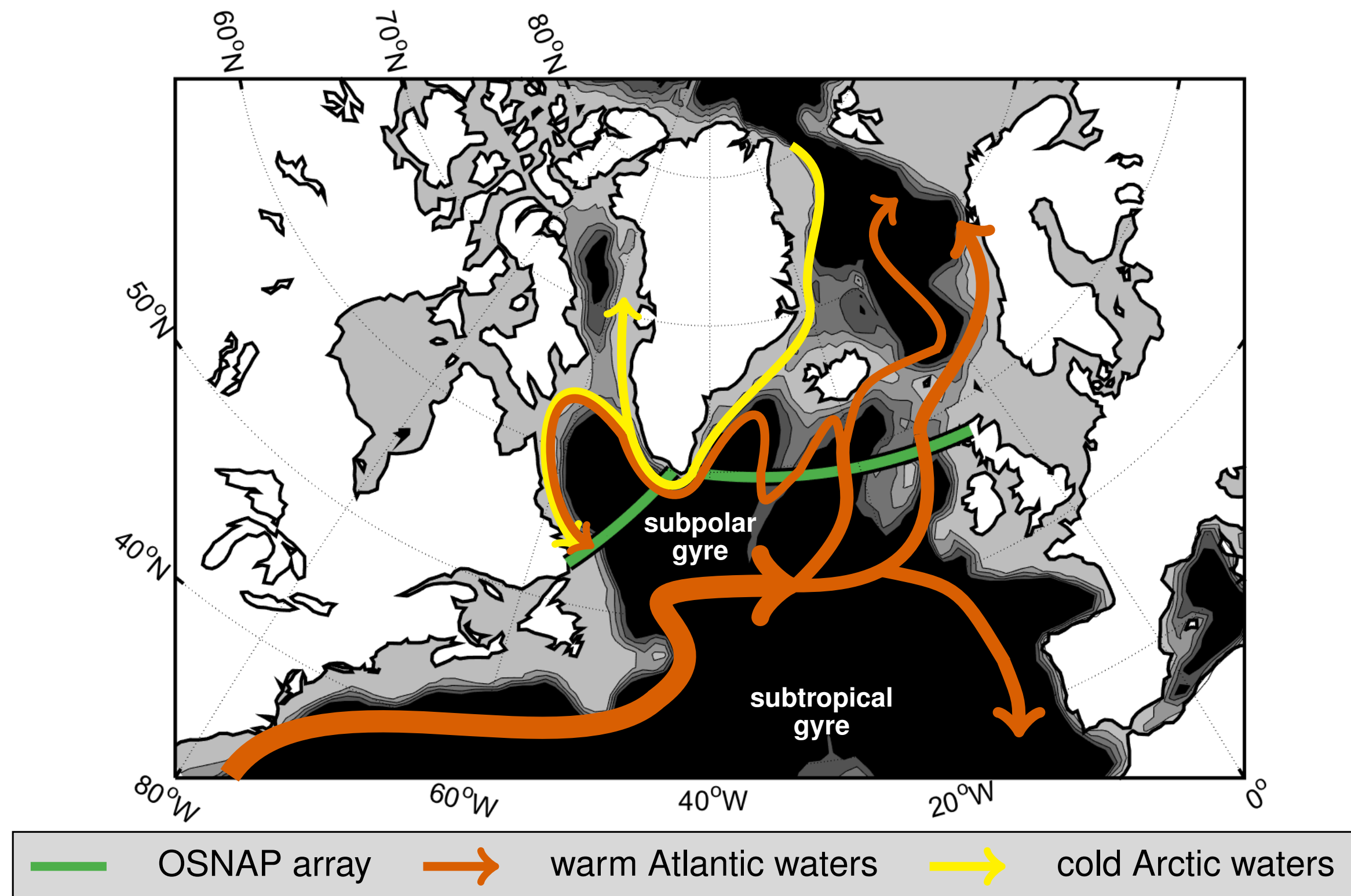
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Algorithmic approach

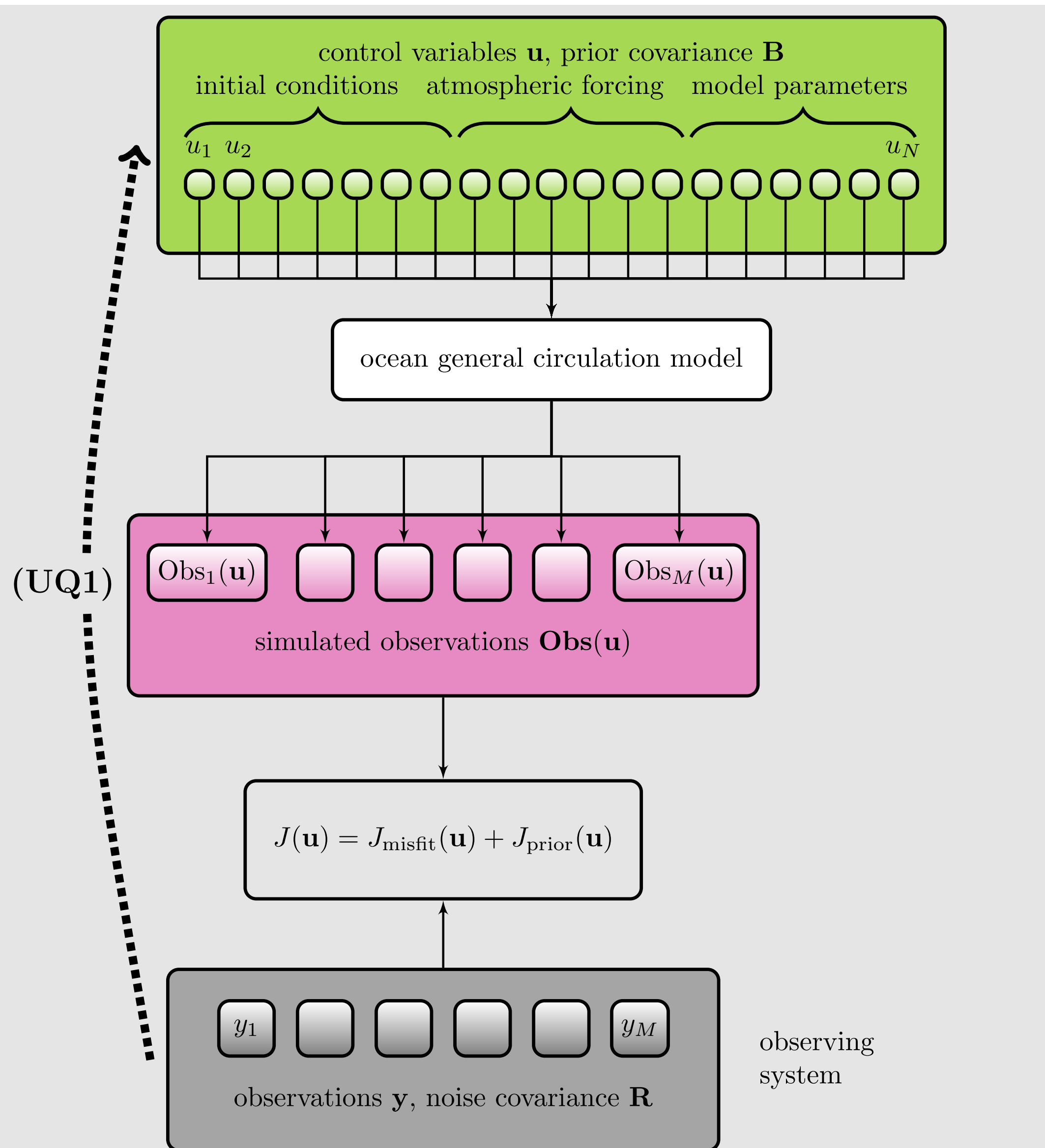
Formulate
Quantity of Interest (QoI)

Leverage adjoint-based
data assimilation
framework:
ECCO



Uncertainty Quantification (UQ)

(UQ1): Inverse uncertainty propagation
From prior to posterior uncertainty



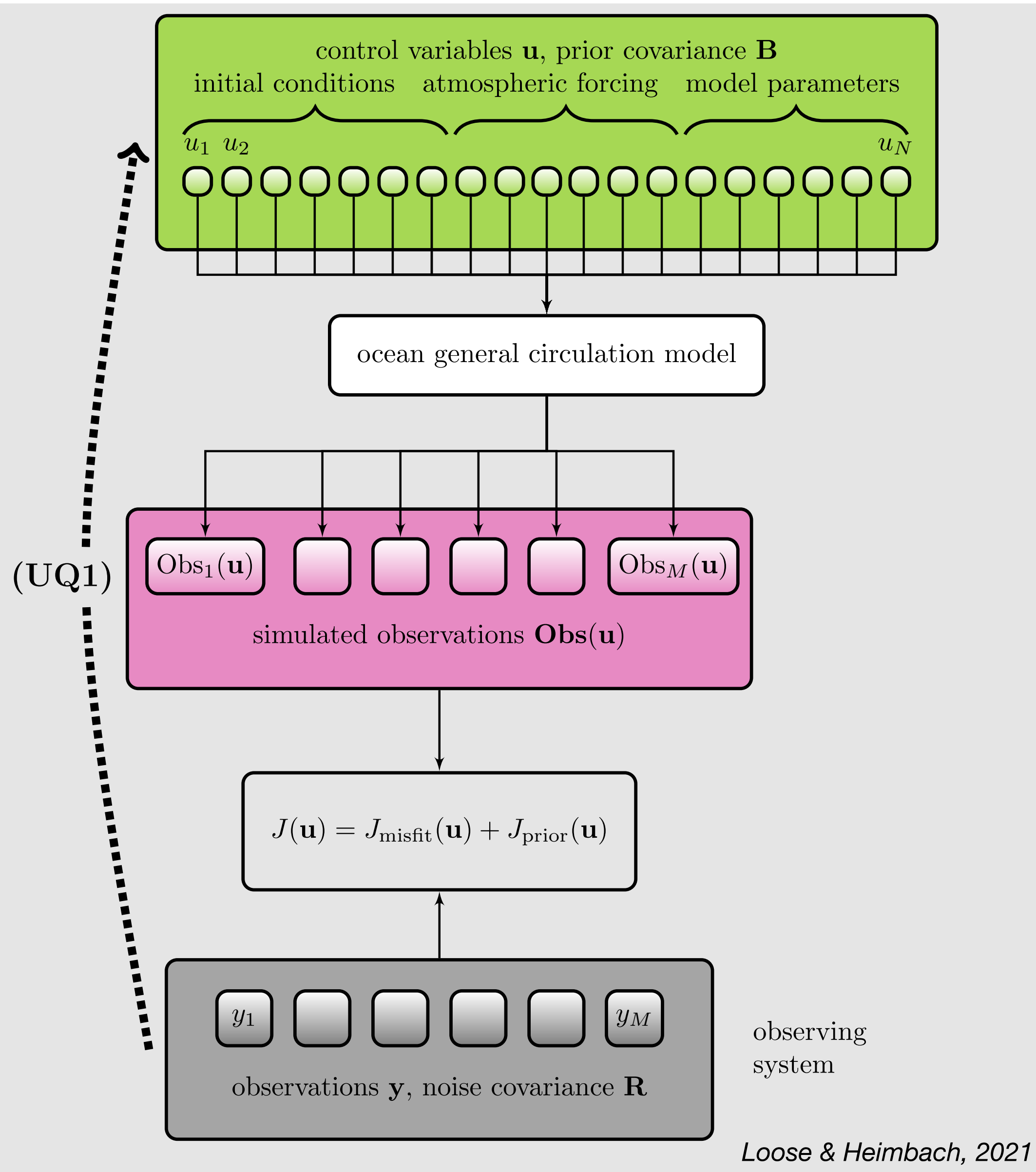
$$J(\mathbf{u}) = \frac{1}{2} (\mathbf{y} - \text{Obs}(\mathbf{u}))^T \mathbf{R}^{-1} (\mathbf{y} - \text{Obs}(\mathbf{u})) + \frac{1}{2} (\mathbf{u} - \mathbf{u}_0)^T \mathbf{B}^{-1} (\mathbf{u} - \mathbf{u}_0).$$

The equation is annotated with arrows indicating the source of each term:

- \mathbf{y} (Observations) and $\text{Obs}(\mathbf{u})$ (Simulated observations) contribute to the misfit term $J_{\text{misfit}}(\mathbf{u})$.
- \mathbf{u}_0 (First guess) and \mathbf{B} (Prior covariance) contribute to the prior term $J_{\text{prior}}(\mathbf{u})$.
- Labels below the equation identify the uncertainty sources: "Uncertain model inputs" points to \mathbf{u} , "Observation uncertainties" points to \mathbf{R} , and "Prior uncertainties" points to \mathbf{B} .

Uncertainty Quantification (UQ)

(UQ1): Inverse uncertainty propagation
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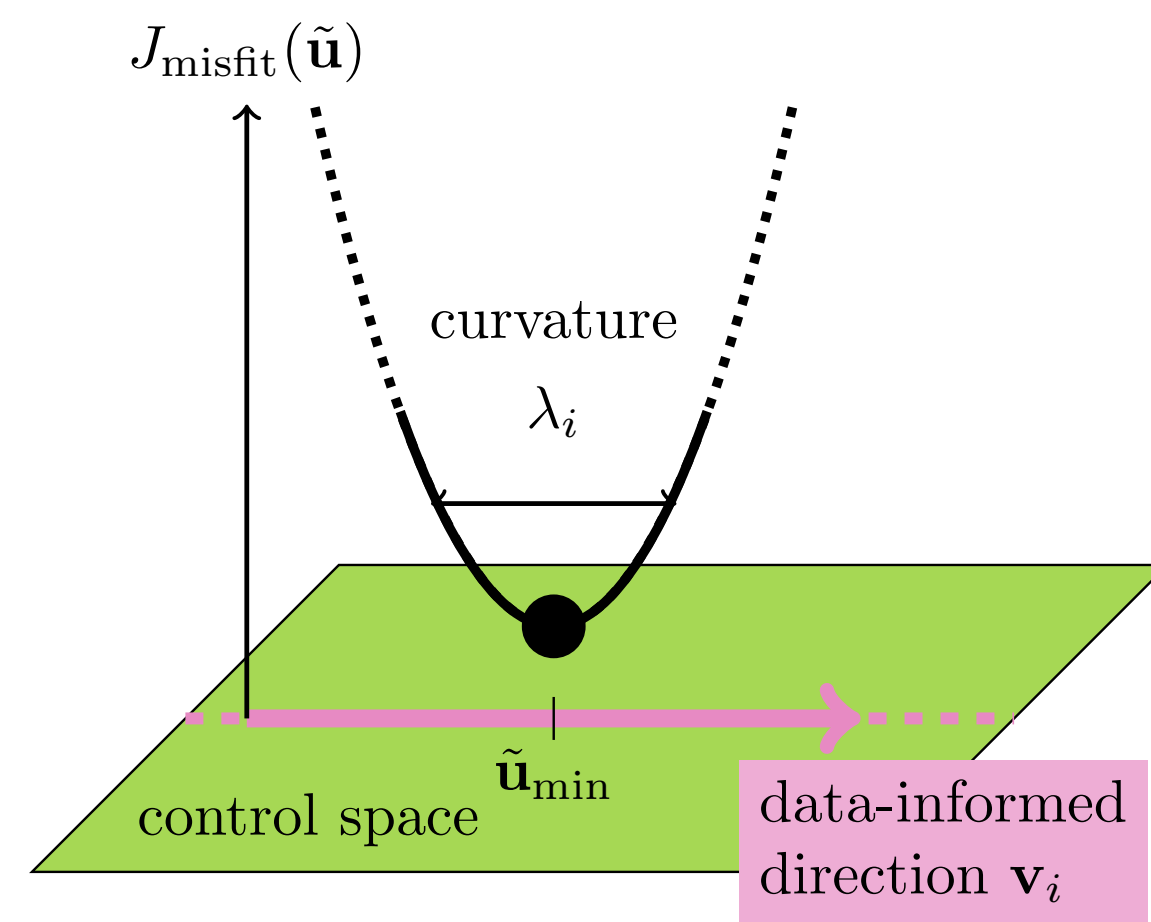


Observations Simulated observations First guess

$$J(\mathbf{u}) = \frac{1}{2} (\mathbf{y} - \text{Obs}(\mathbf{u}))^T \mathbf{R}^{-1} (\mathbf{y} - \text{Obs}(\mathbf{u})) + \frac{1}{2} (\mathbf{u} - \mathbf{u}_0)^T \mathbf{B}^{-1} (\mathbf{u} - \mathbf{u}_0)$$

$J_{\text{misfit}}(\mathbf{u})$ $J_{\text{prior}}(\mathbf{u})$

Uncertain model inputs Observation uncertainties Prior uncertainties



Hessian of cost function J

$$\mathbf{H}_J \approx \sum_{i=1}^M \lambda_i \mathbf{v}_i \mathbf{v}_i^T$$

Thacker, 1989

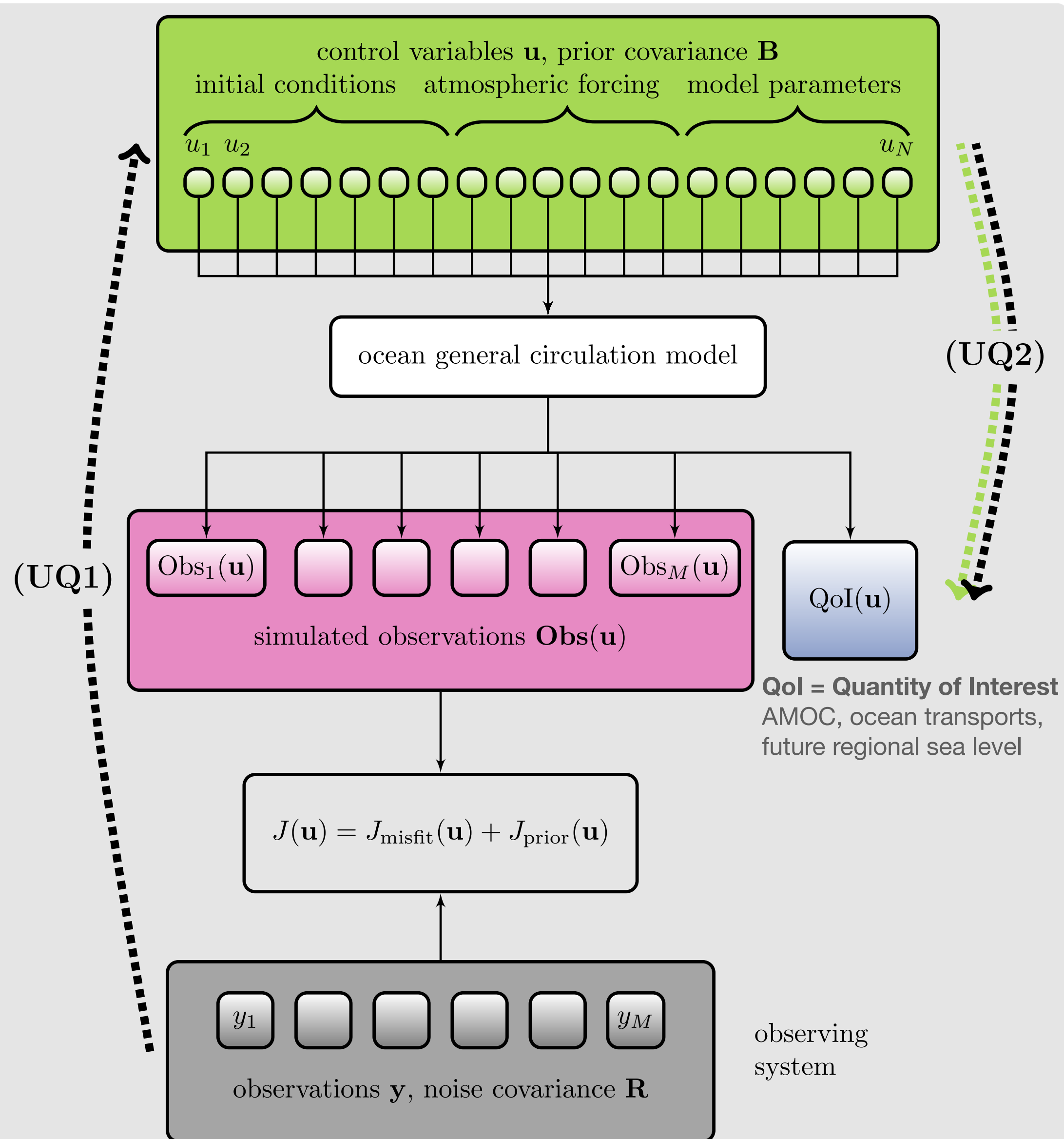
posterior uncertainty ~ inverse Hessian of J

Uncertainty Quantification (UQ)

(UQ2): Forward uncertainty propagation From posterior to QoI uncertainty

Uncertainty reduction in quantity of interest (QoI) by observing system:

$$\sum_{i=1}^M \frac{\lambda_i}{\lambda_i + 1} (\mathbf{q} \cdot \mathbf{v}_i)^2 \in [0, 1) \quad \text{where} \quad \mathbf{q} = \frac{\mathbf{B}^{T/2} \nabla_u \text{QoI}}{\|\mathbf{B}^{T/2} \nabla_u \text{QoI}\|}$$



Loose & Heimbach, 2021

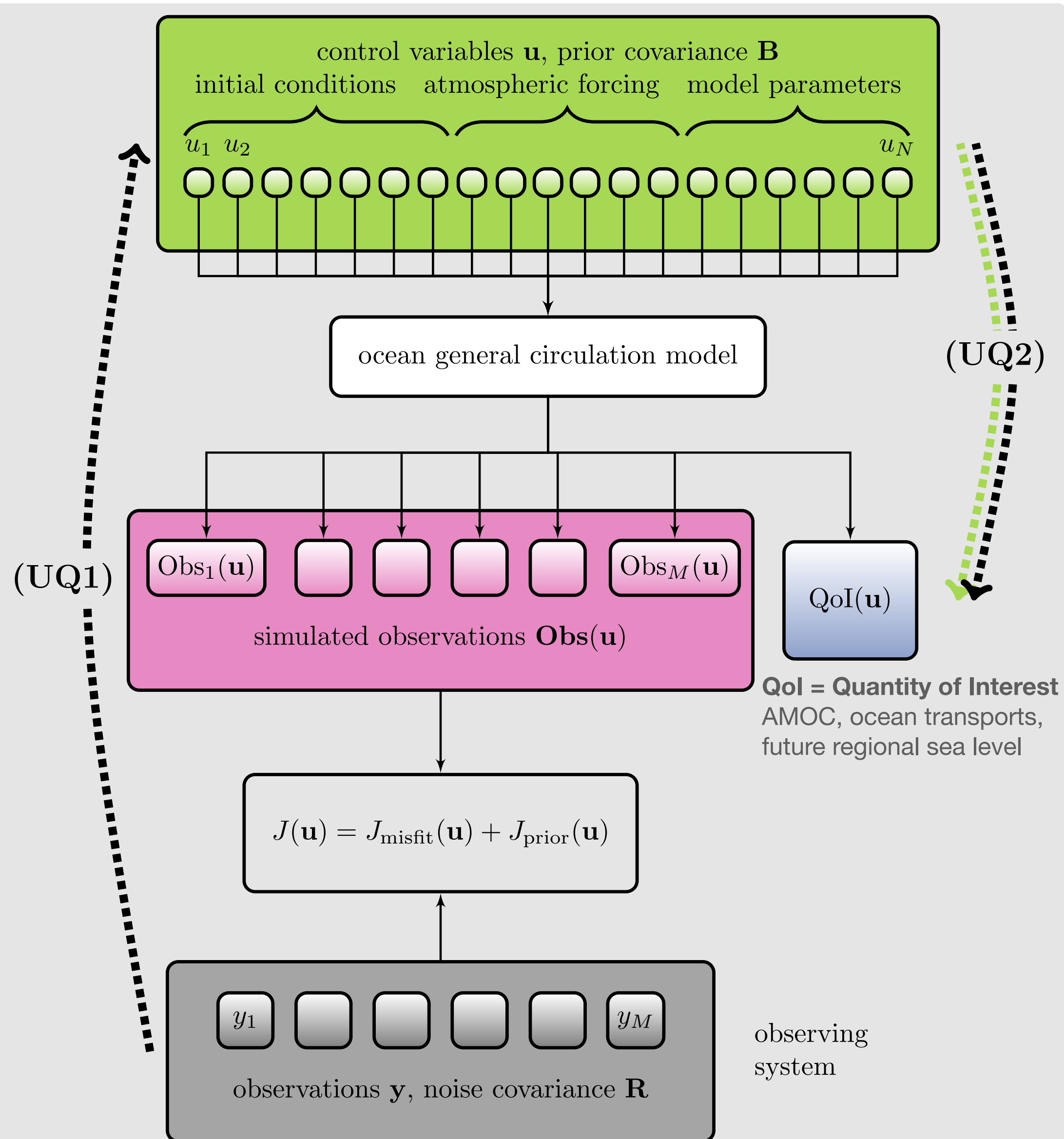
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$(\mathbf{q} \cdot \mathbf{v}_i)^2 =$ Degree of shared adjustment mechanisms between observing system and QoI



Loose & Heimbach, 2021

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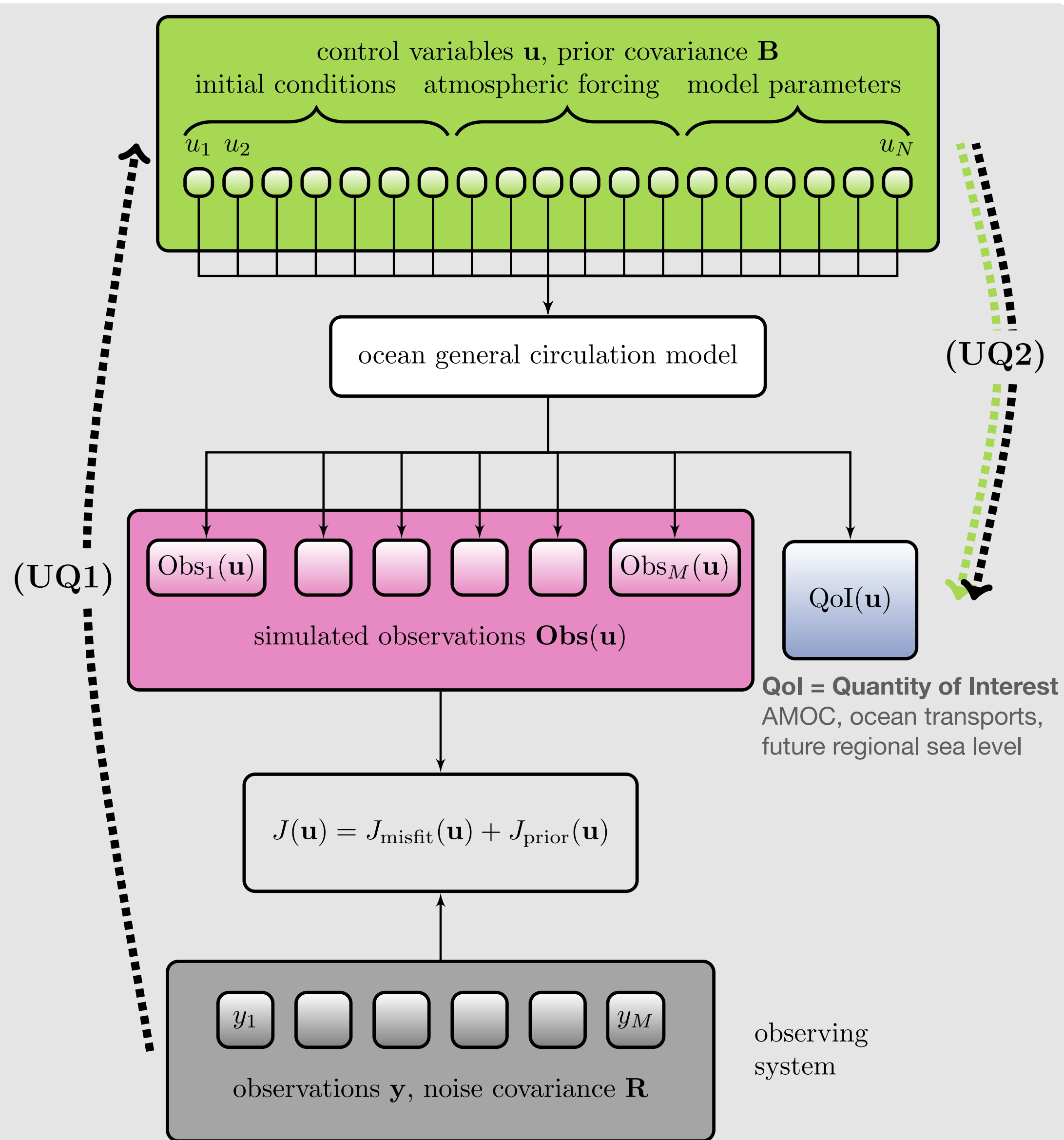
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Eigenvalues reflect "sensitivity-to-noise" ratio of observations

$$\lambda = \frac{\|\mathbf{B}^{T/2} \nabla_u \text{Obs}\|^2}{\varepsilon_{\text{Obs}}^2}$$



Loose & Heimbach, 2021

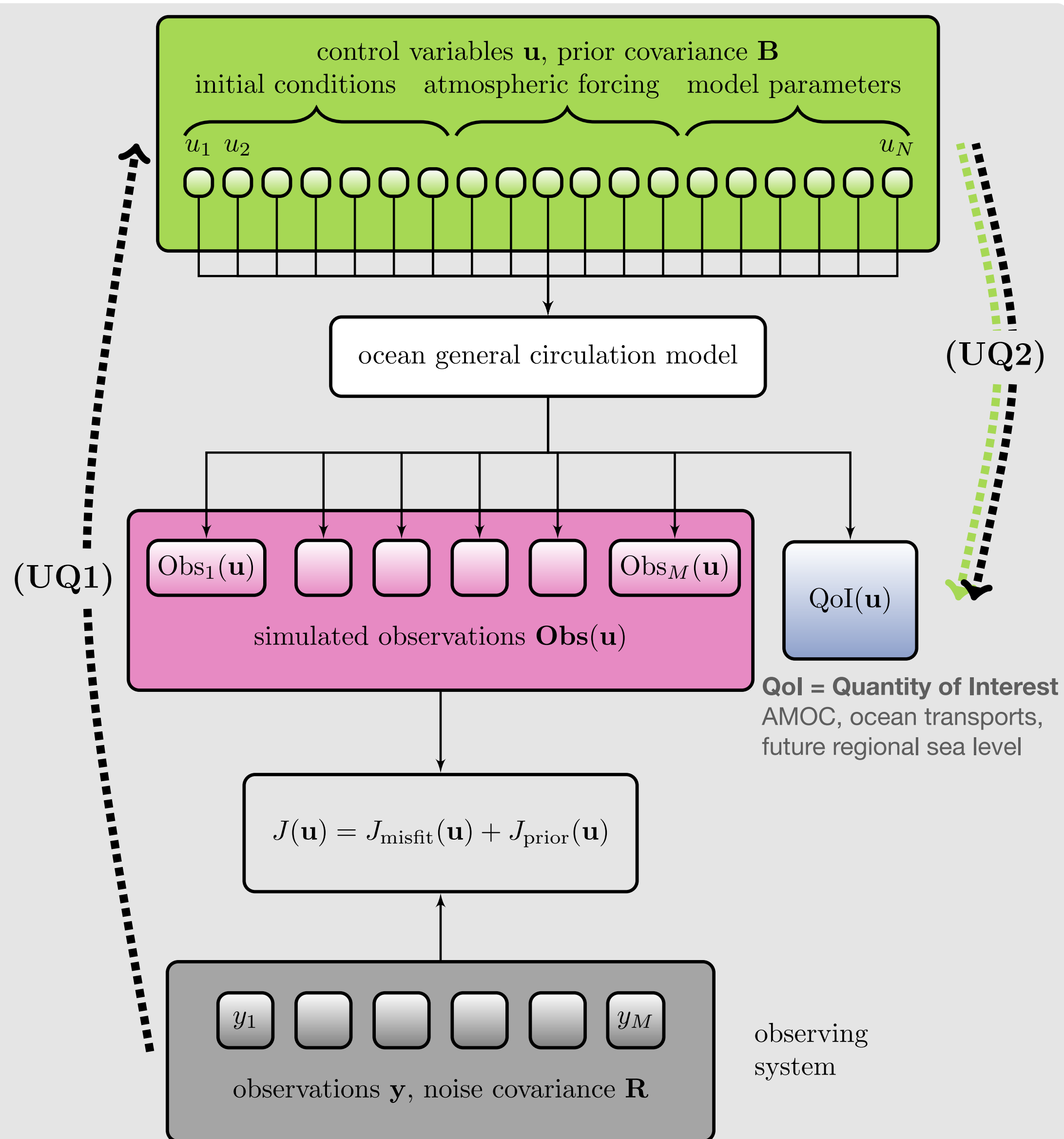
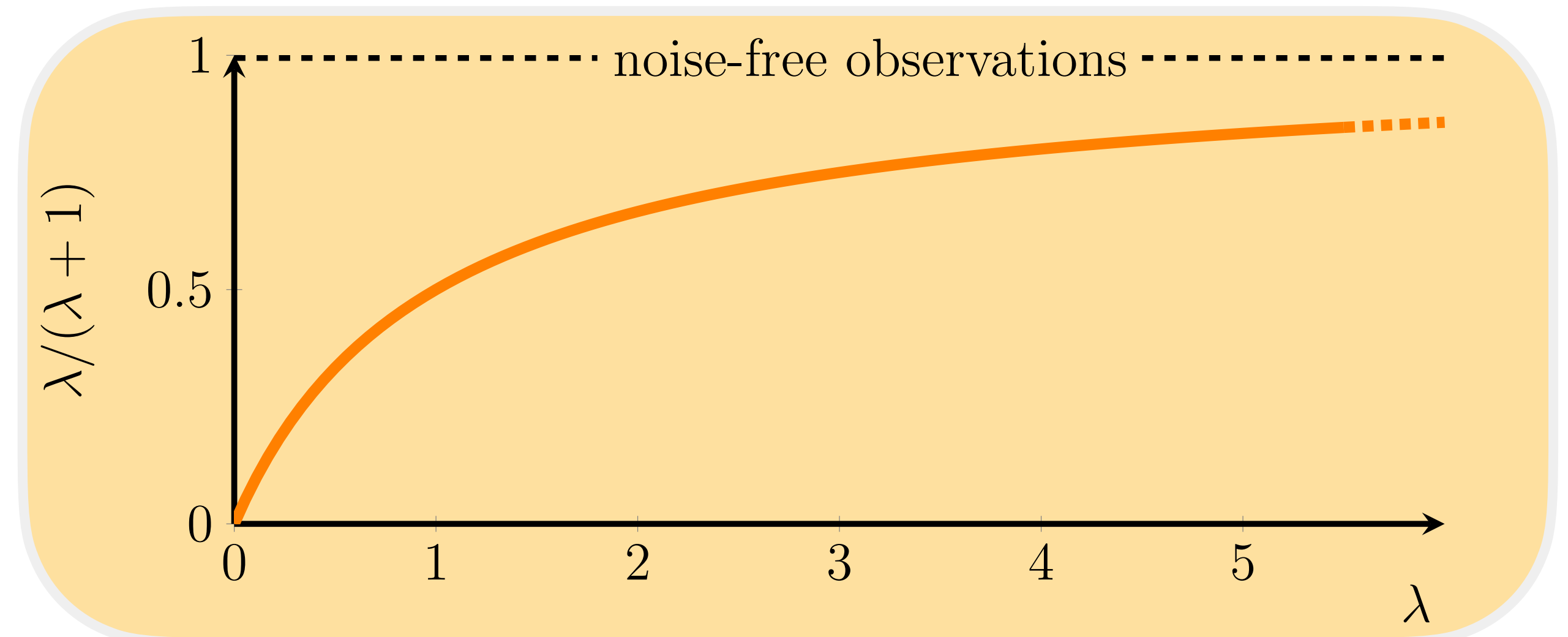
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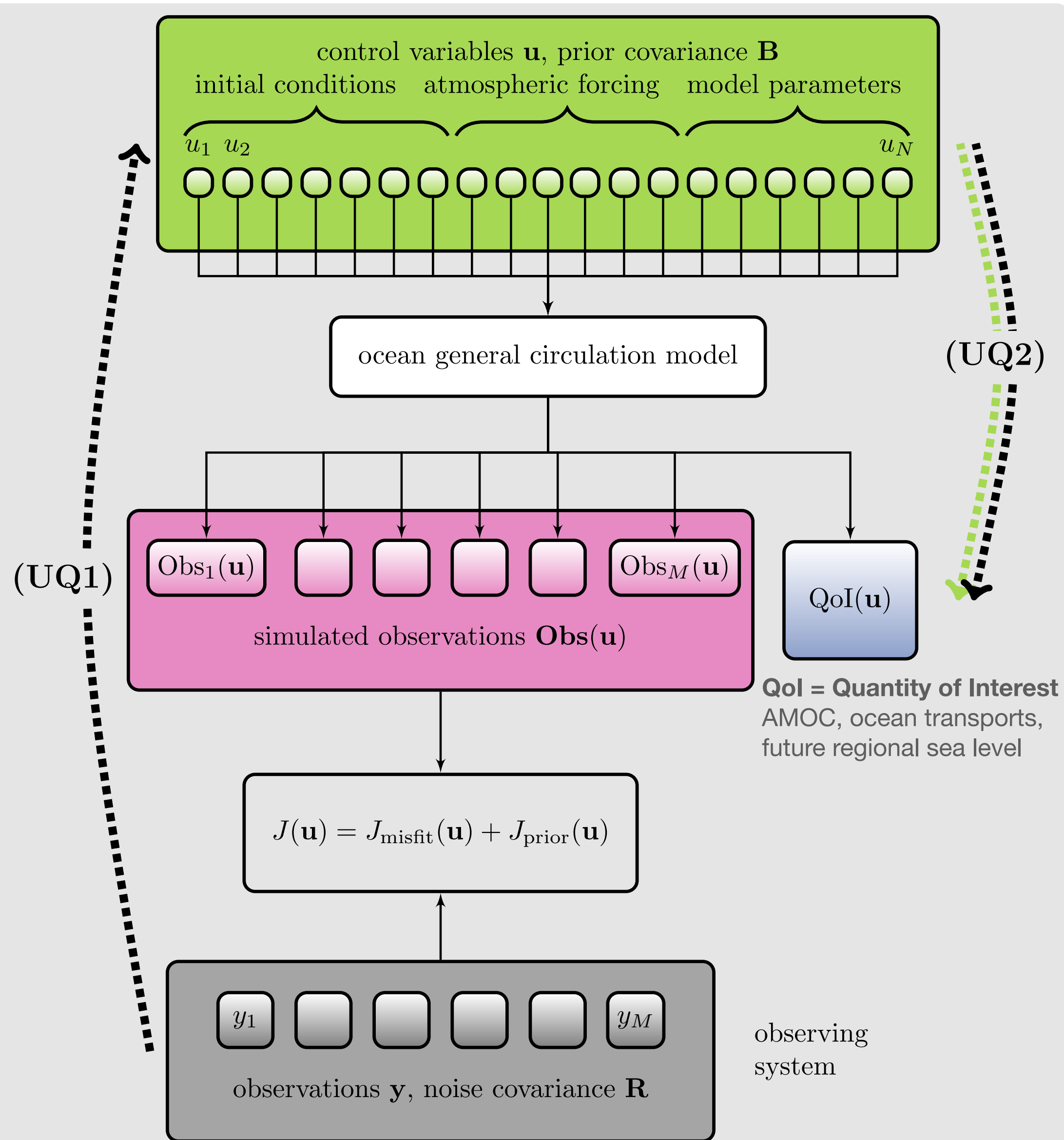
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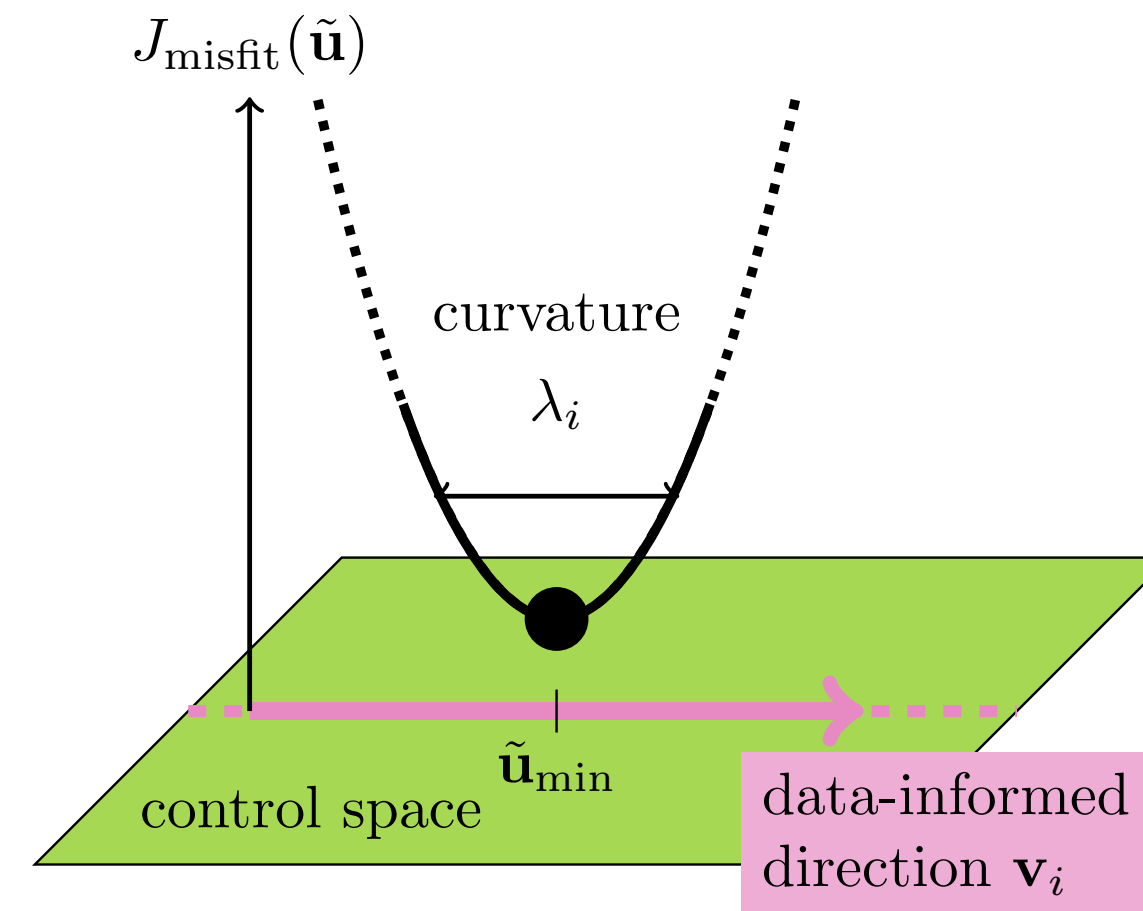


Uncertainty Quantification (UQ)



Loose & Heimbach, 2021

(UQ1): Inverse uncertainty propagation



Hessian of cost function J

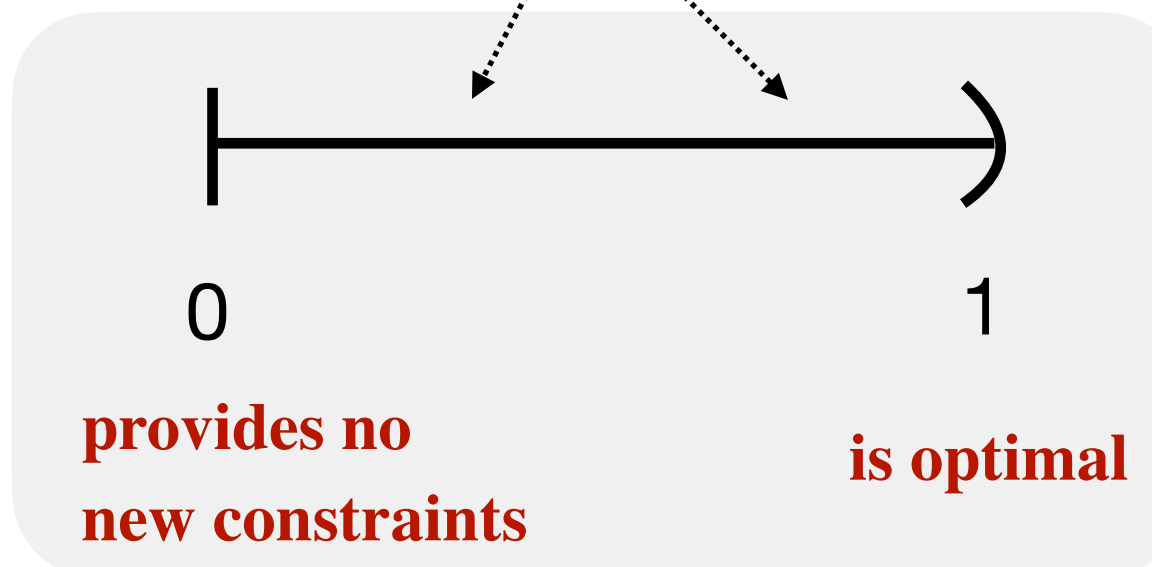
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Thacker, 1989

(UQ2): Forward uncertainty propagation

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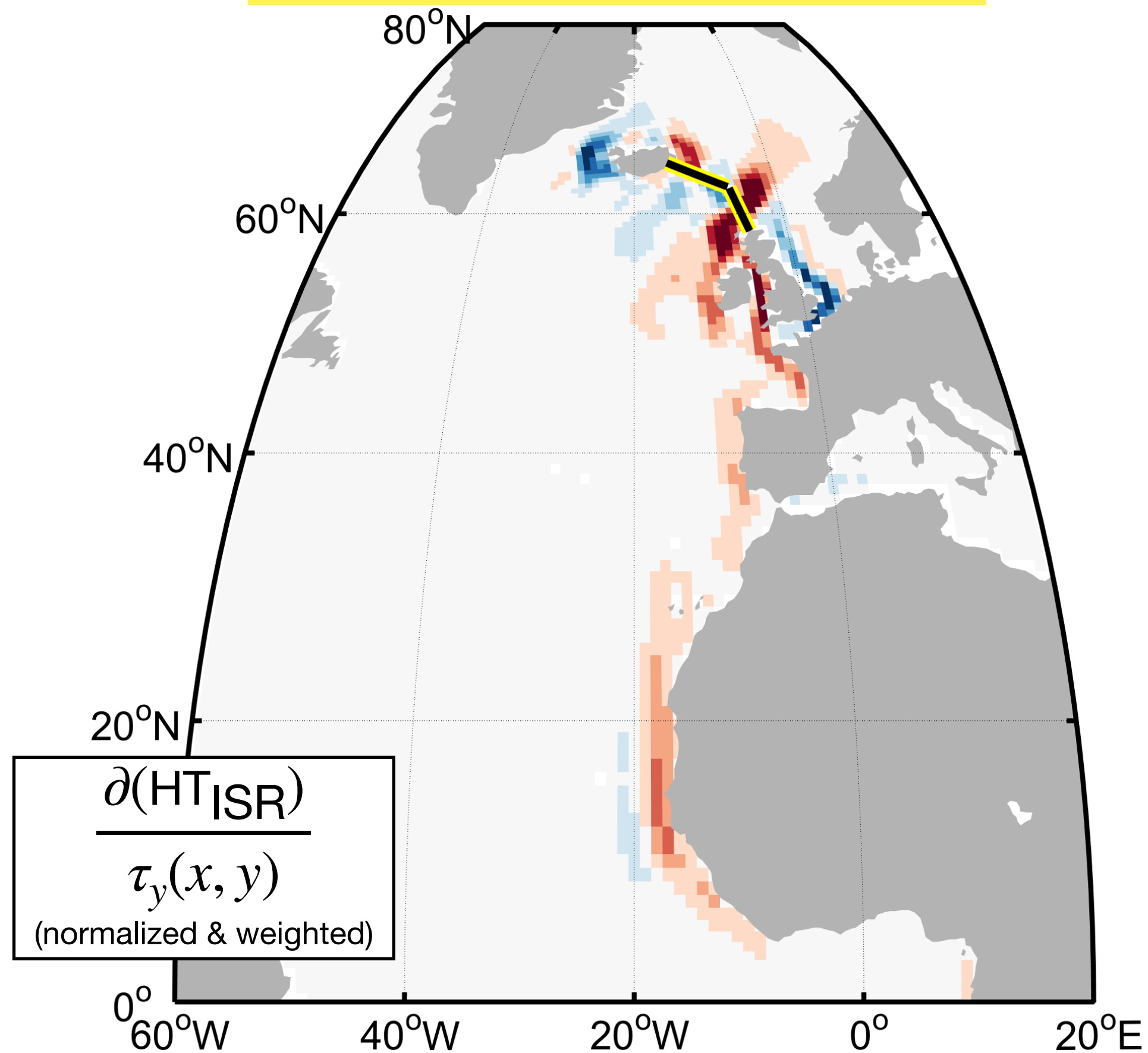


Observing system

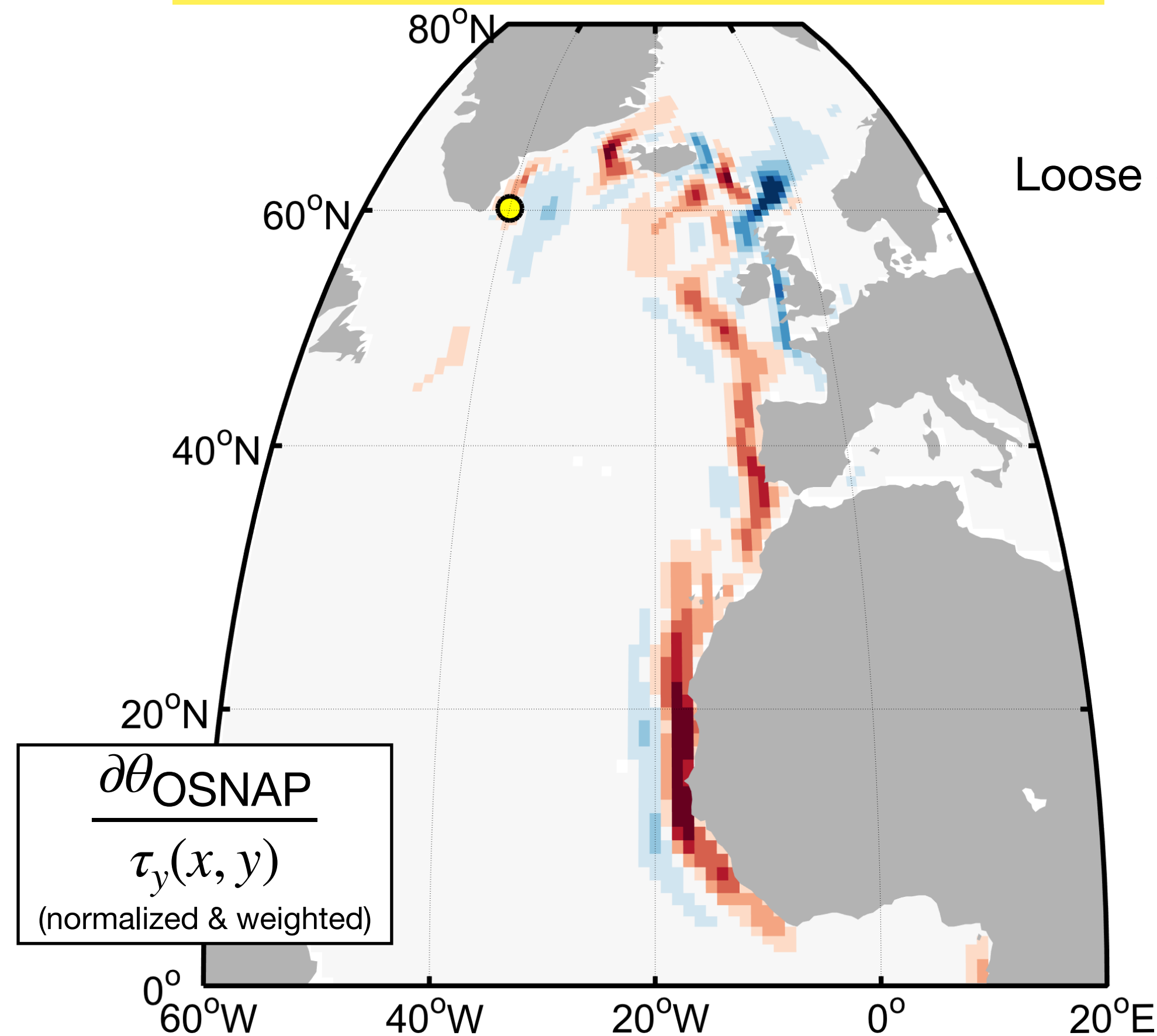
for QoI chosen

Sensitivity maps identify shared adjustment mechanisms & pathways

Heat transport across the Iceland-Scotland Ridge (ISR)

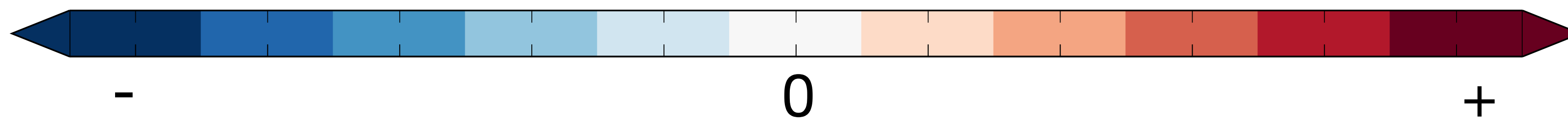


Irminger Sea subsurface temperature (observed by OSNAP)



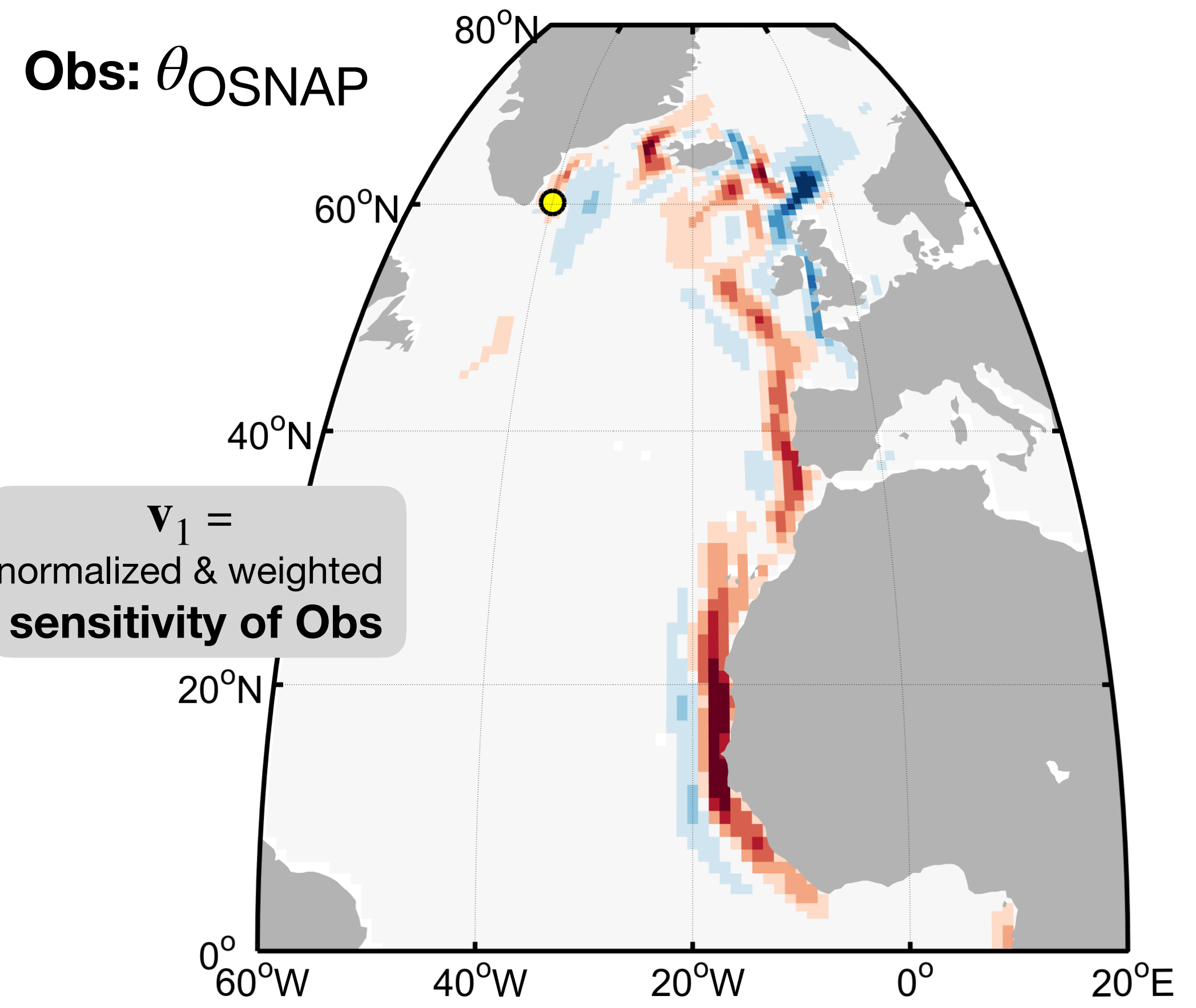
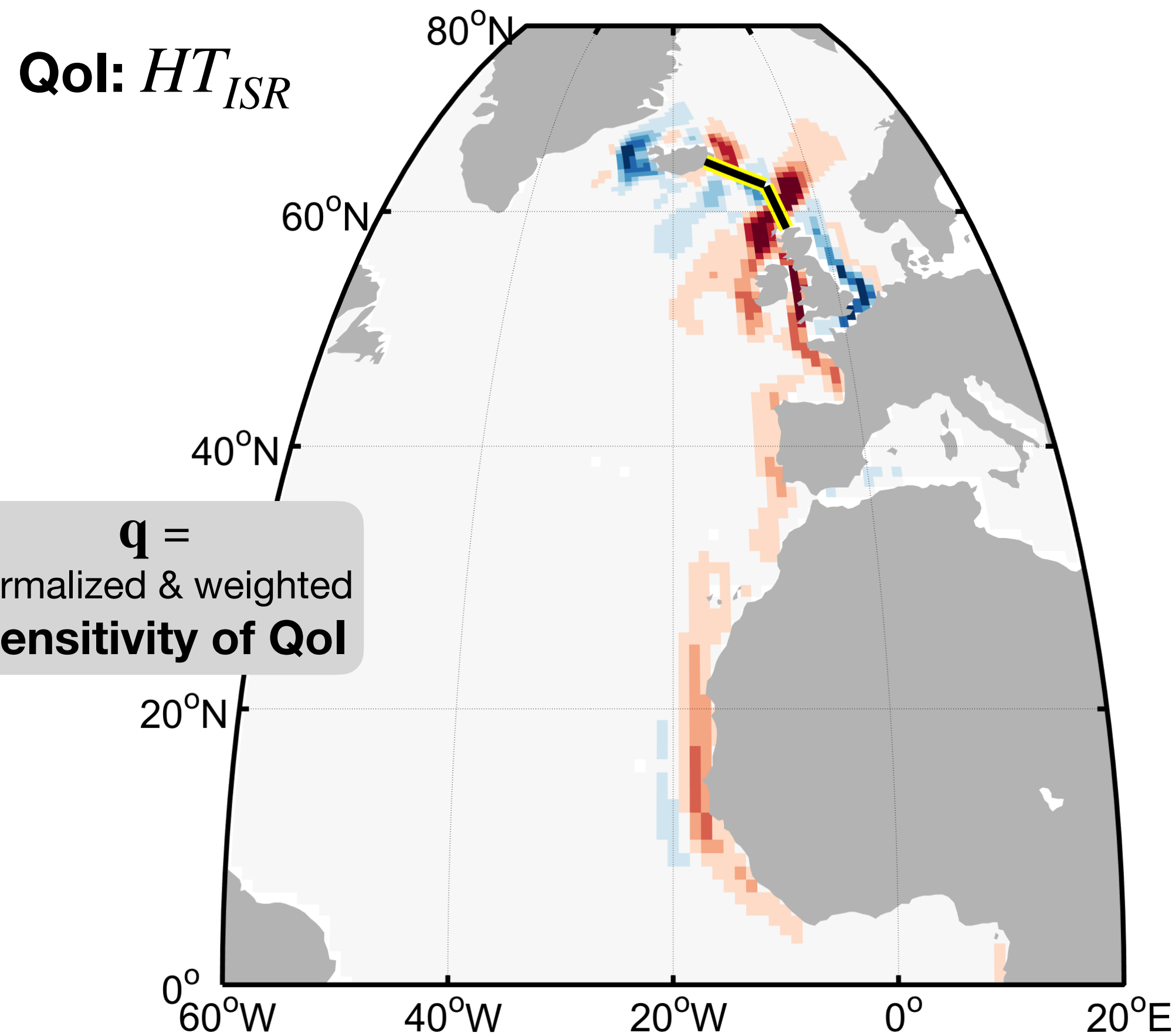
Loose et al., 2020

Sensitivity to meridional wind stress τ_y



Computed with MITgcm adjoint model

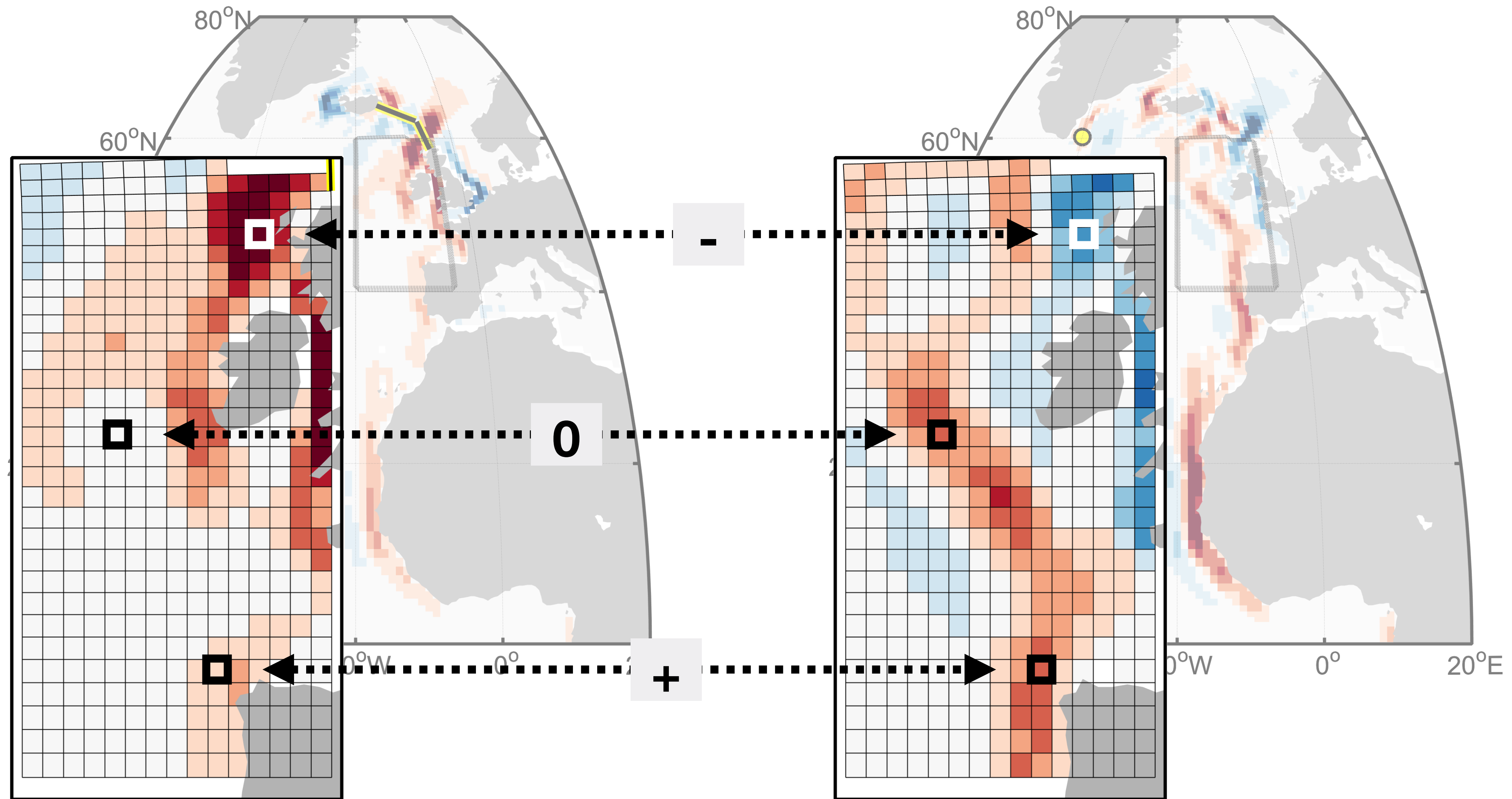
Sensitivity maps re-interpreted in the context of UQ



v₁ = Hessian eigenvector (in ECCO)

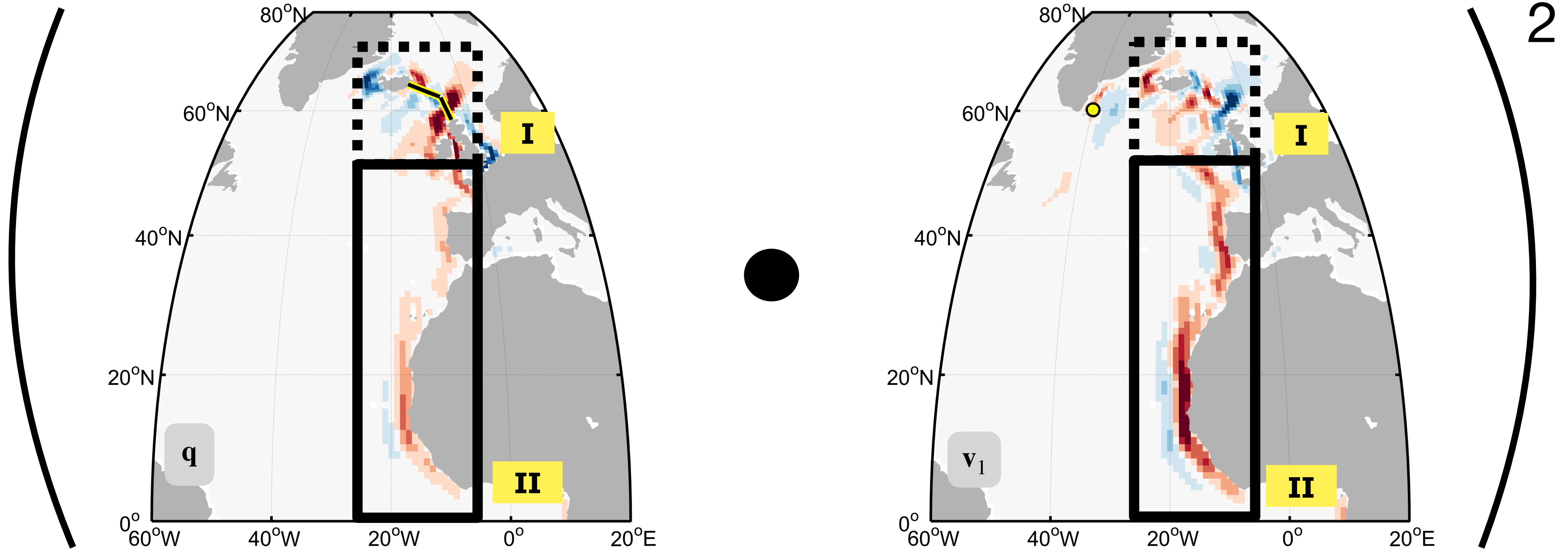
$(\mathbf{q} \bullet \mathbf{v}_1)^2$ = uncertainty reduction in HT_{ISR} by noise-free θ_{OSNAP} observation

Computing $(\mathbf{q} \cdot \mathbf{v}_1)$



$$(\mathbf{q} \cdot \mathbf{v}_1) = ([\mathbf{q} \cdot \mathbf{v}_1]_+ + [\mathbf{q} \cdot \mathbf{v}_1]_- + \cancel{[\mathbf{q} \cdot \mathbf{v}_1]_0})$$

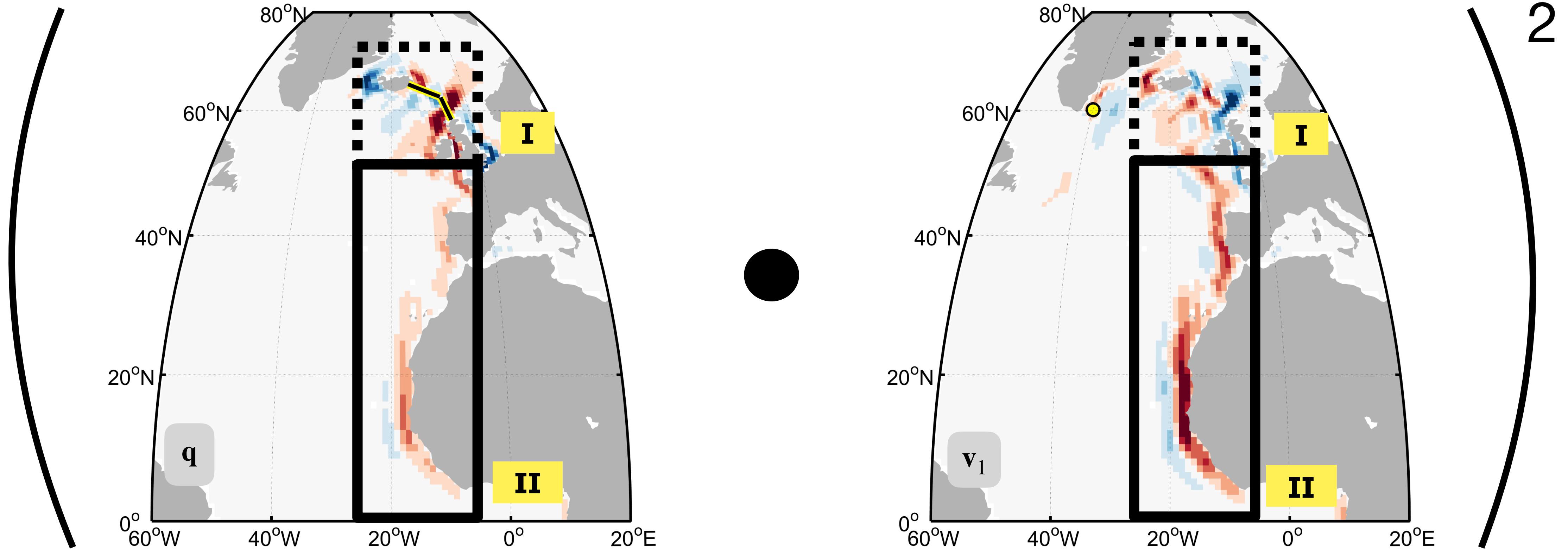
Uncertainty reduction via oceanic teleconnections



Uncertainty reduction in HT_{ISR} by noise-free θ_{OSNAP} observation:

$$(\mathbf{q} \bullet \mathbf{v}_1)^2 = ([\mathbf{q} \bullet \mathbf{v}_1]_{\text{II}} + [\mathbf{q} \bullet \mathbf{v}_1]_{\text{I}})^2 = (0.15 + (-0.59))^2 = (-0.44)^2 = 19\%$$

Uncertainty reduction via oceanic teleconnections

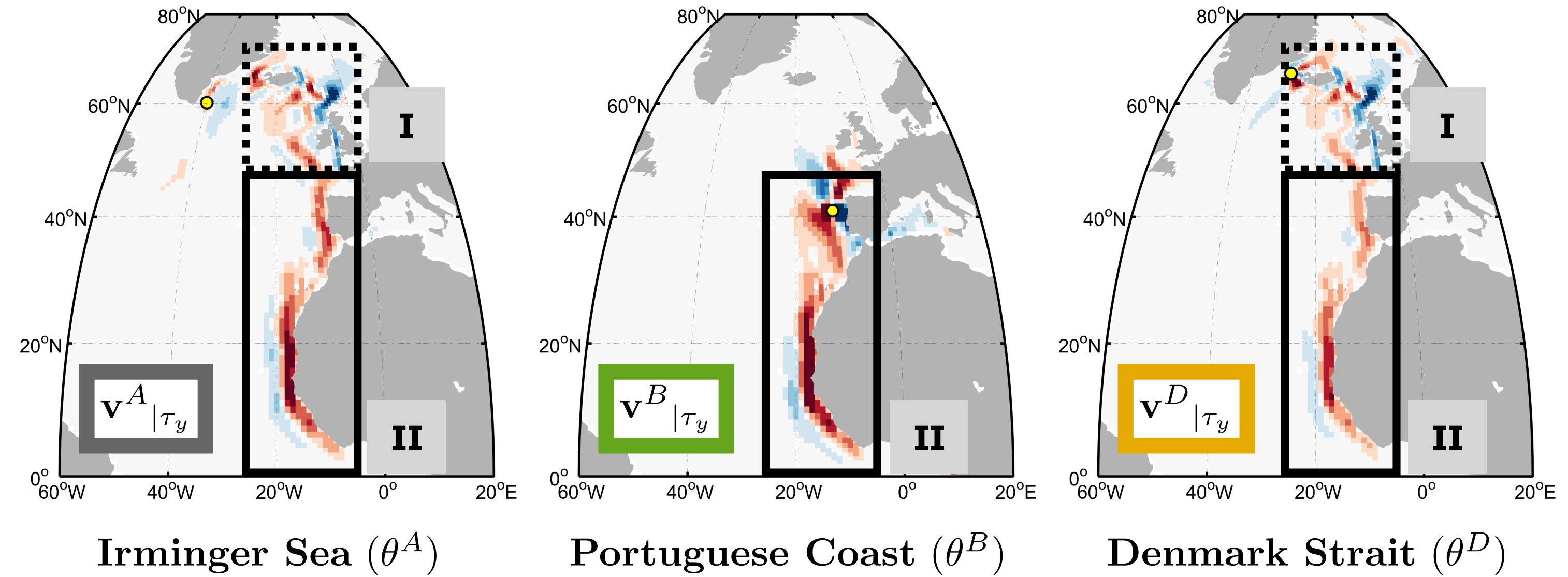


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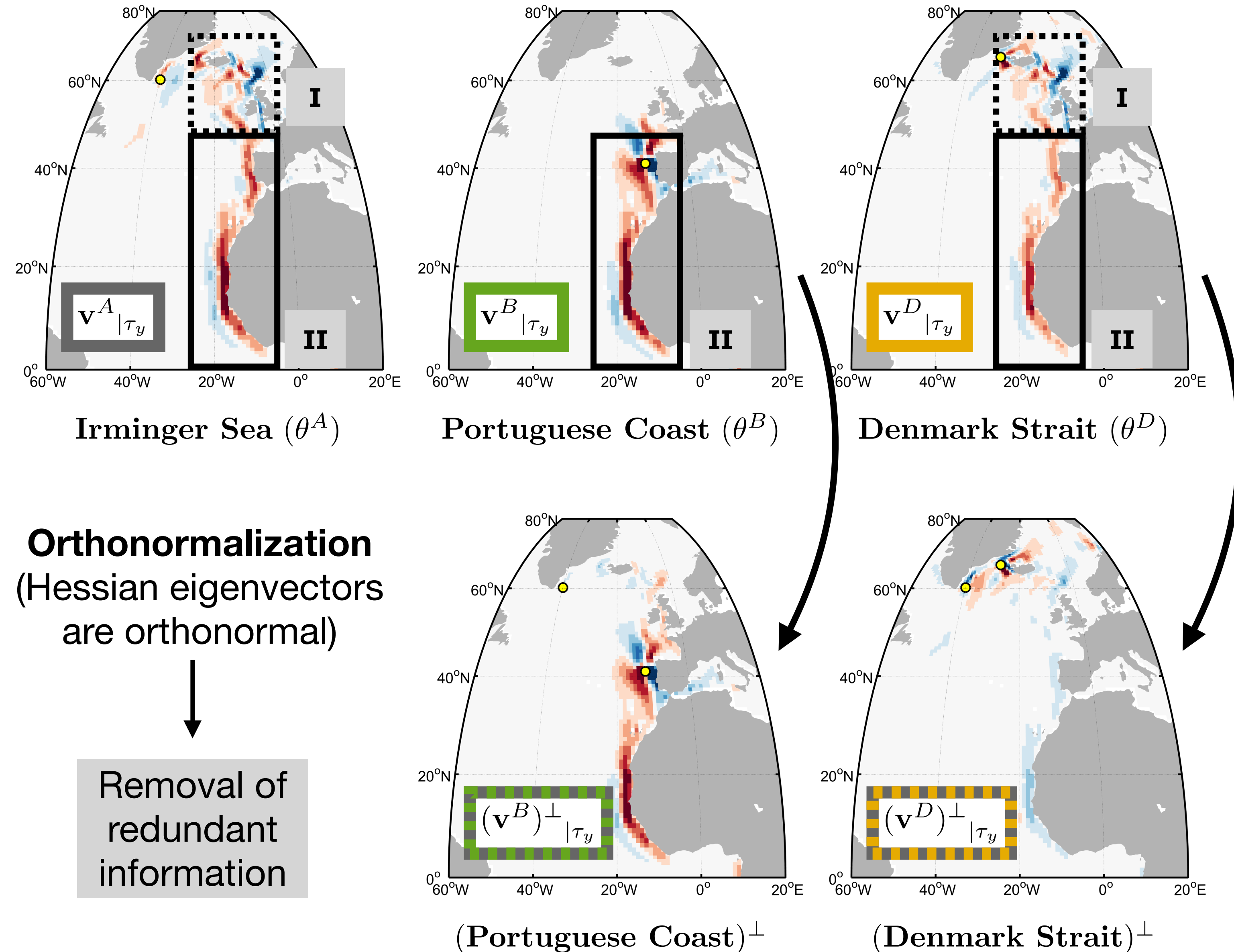
- Shared adjustment mechanisms lead to large uncertainty reduction
- BUT: destructive interference is possible

Assessing data redundancy



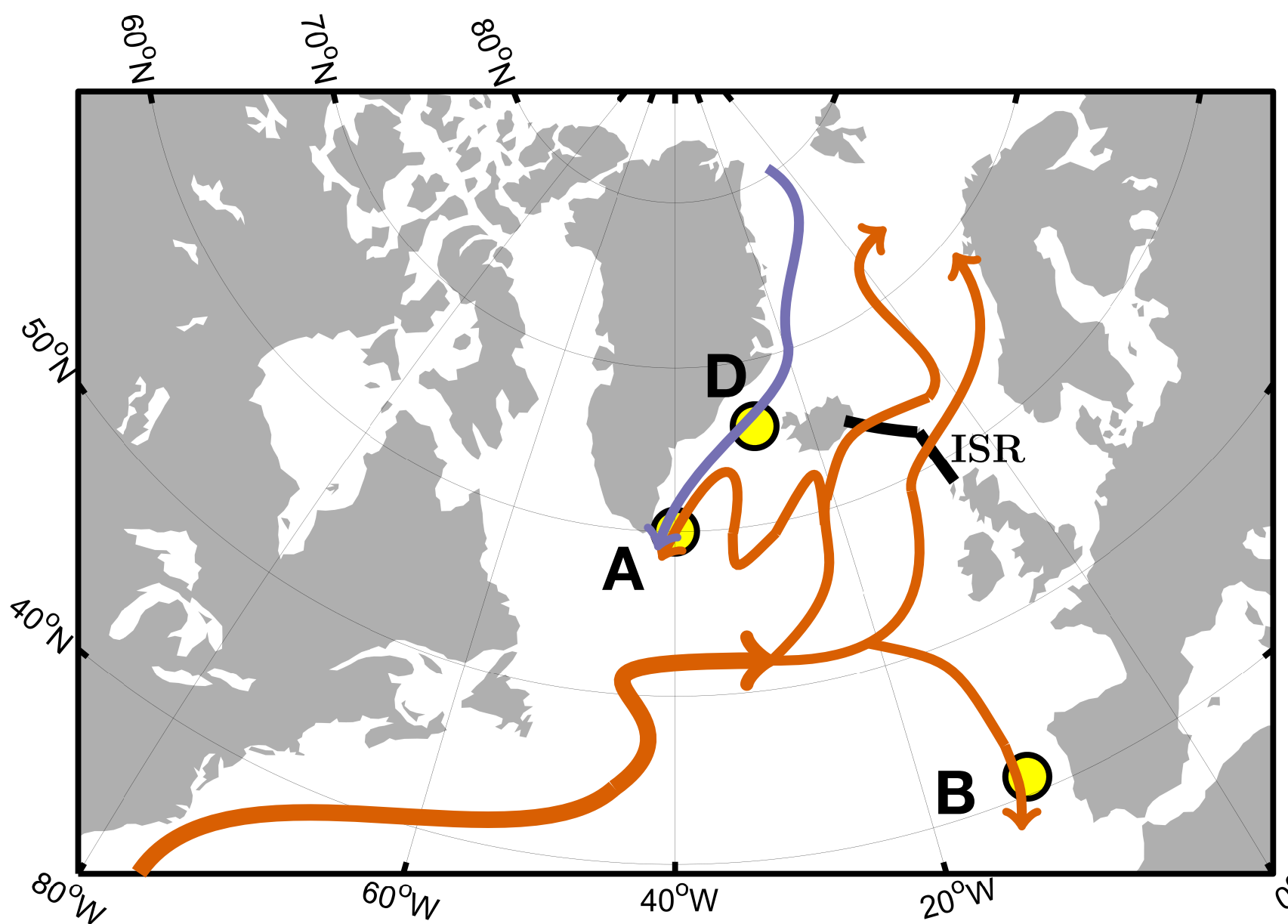
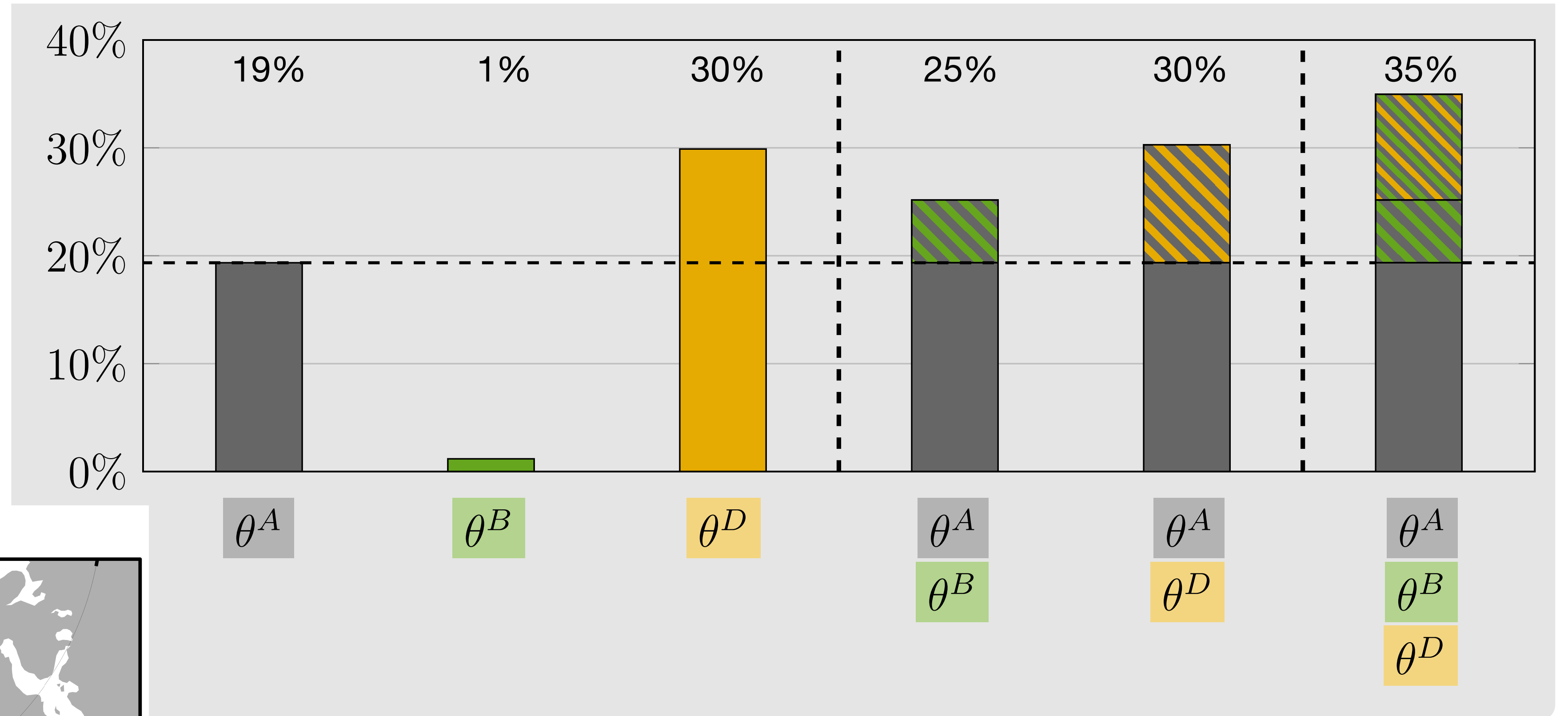
No actual data needed → can test future observing systems

Assessing data redundancy



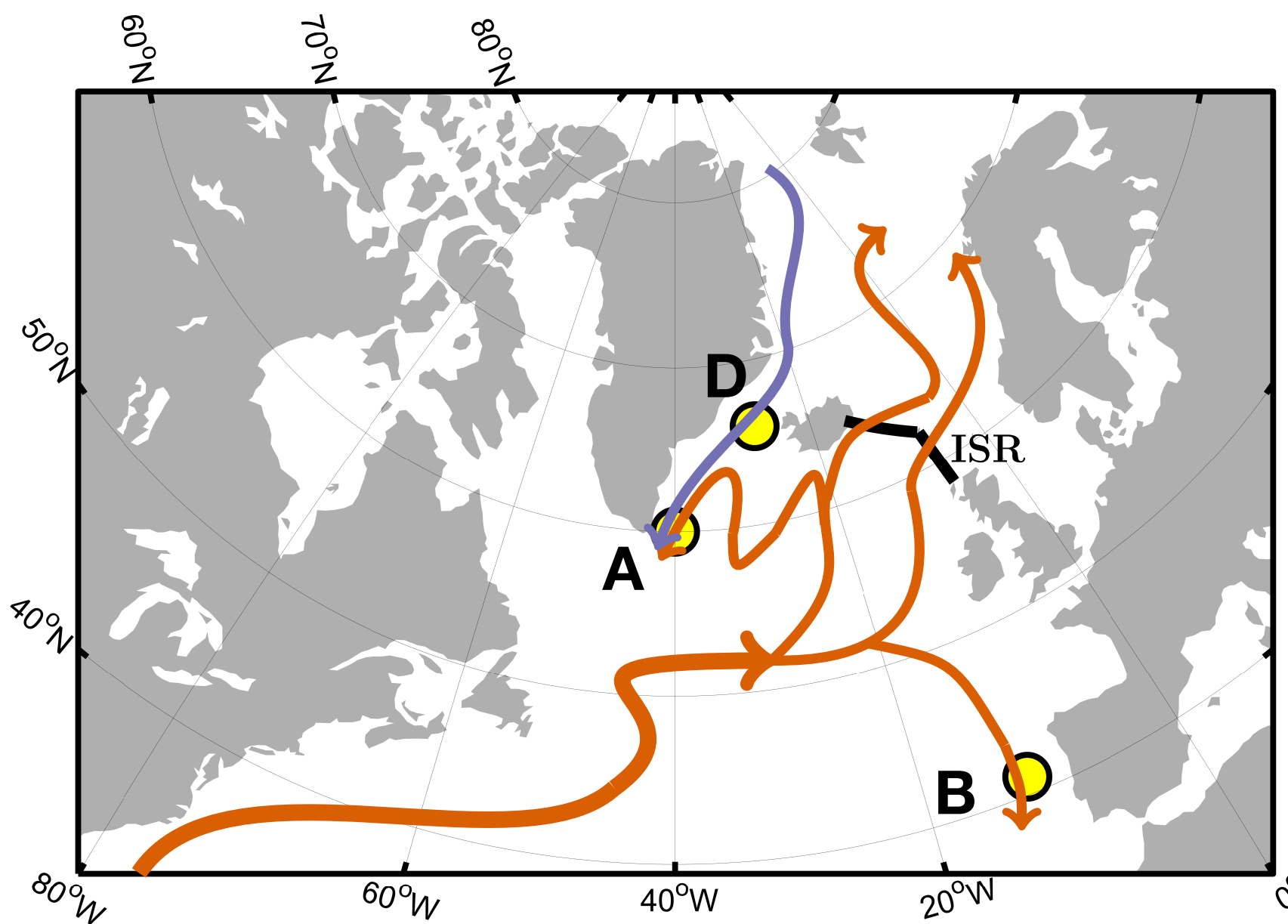
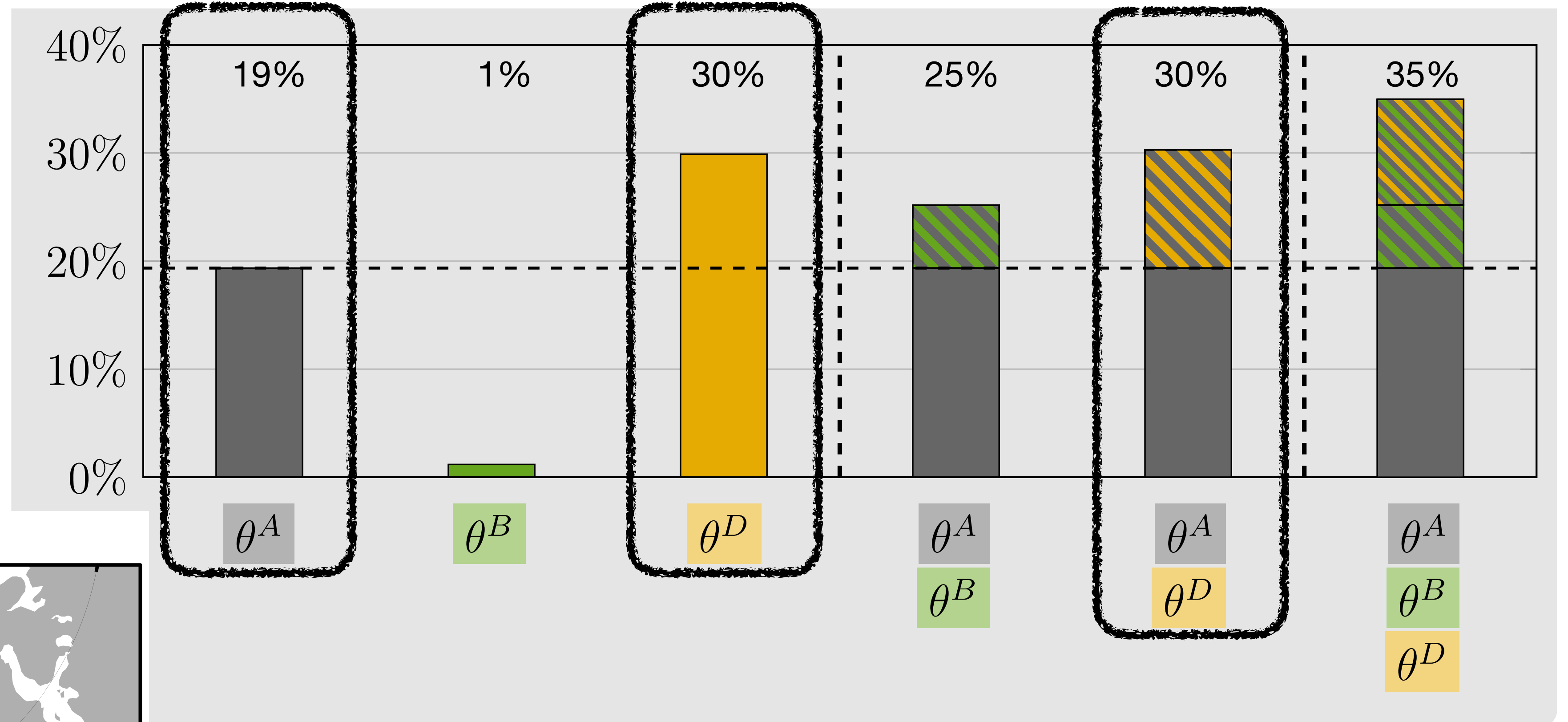
Assessing data redundancy

Uncertainty reduction in HT_{ISR} by noise-free observations



Assessing data redundancy

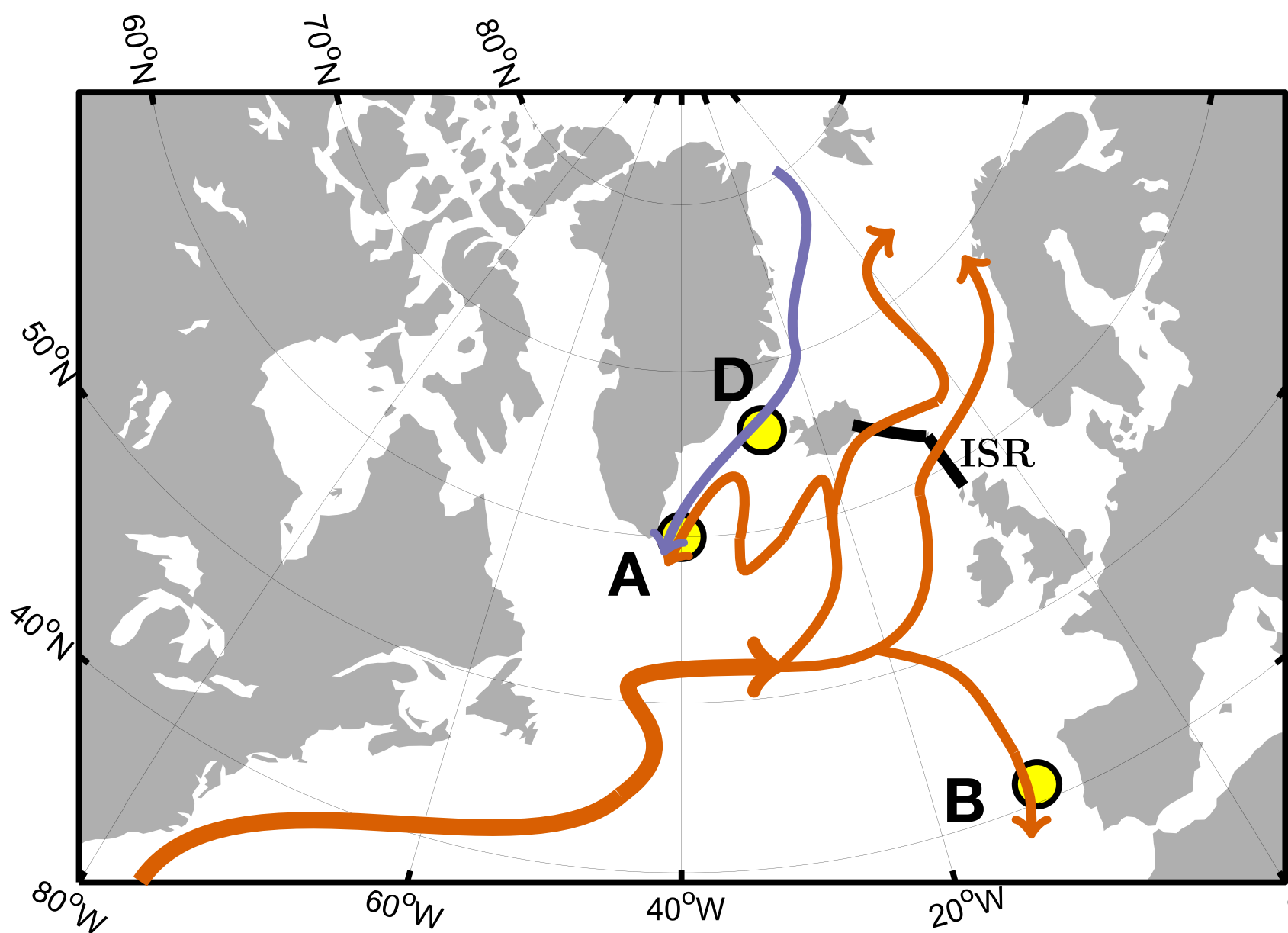
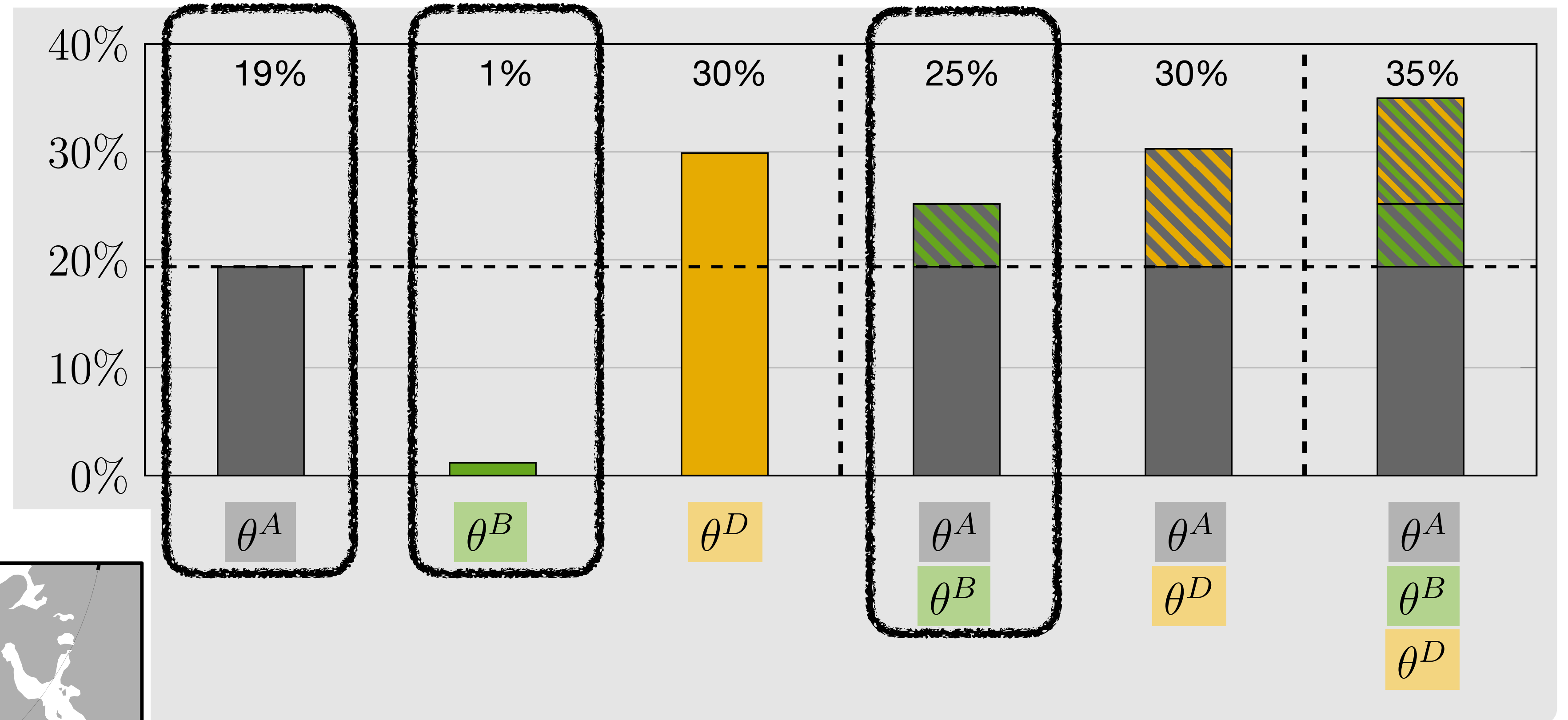
Uncertainty reduction in HT_{ISR} by noise-free observations



- Adding observation A (in the Irminger Sea) to observation D (in Denmark Strait) gives no extra information due to **data redundancy**

Assessing data redundancy

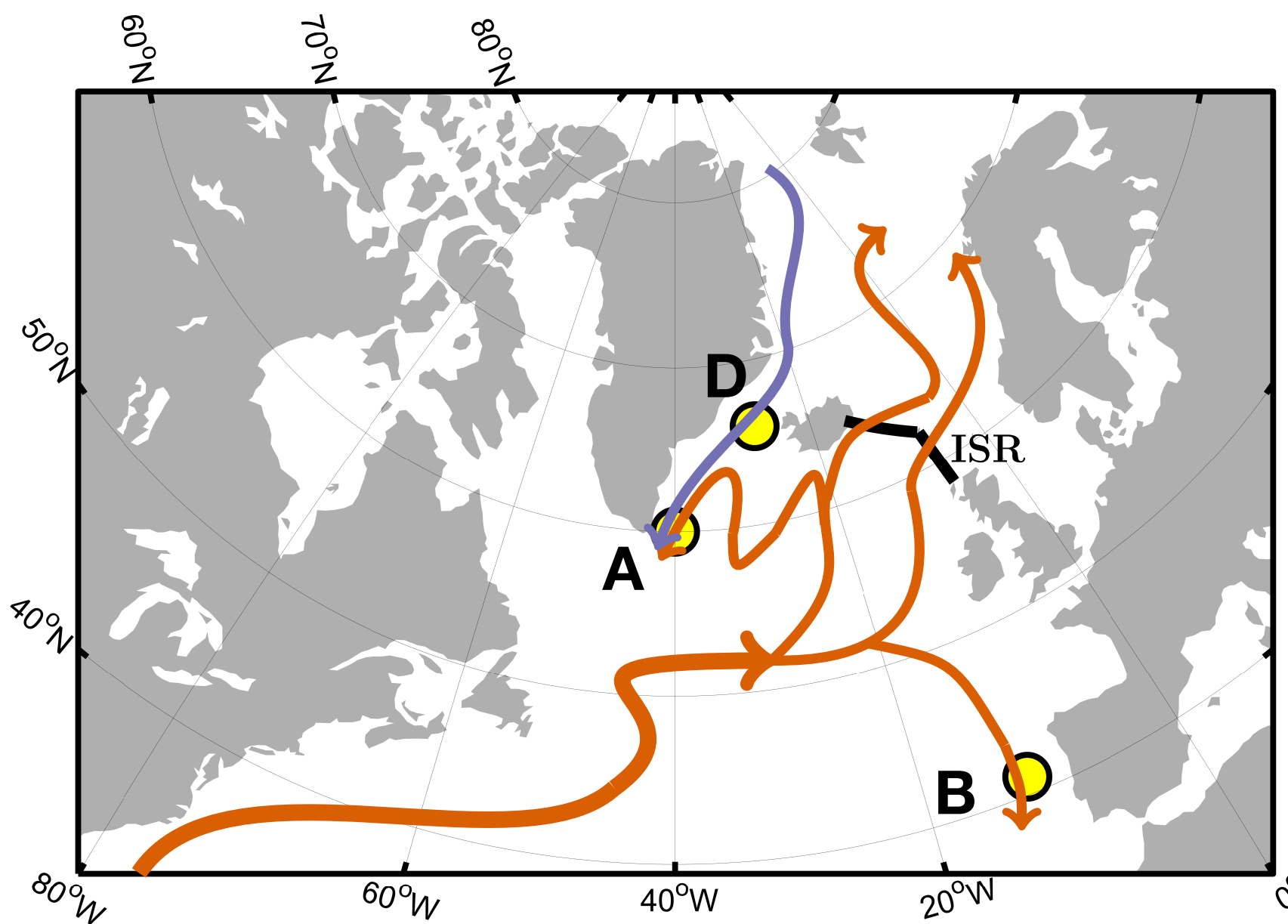
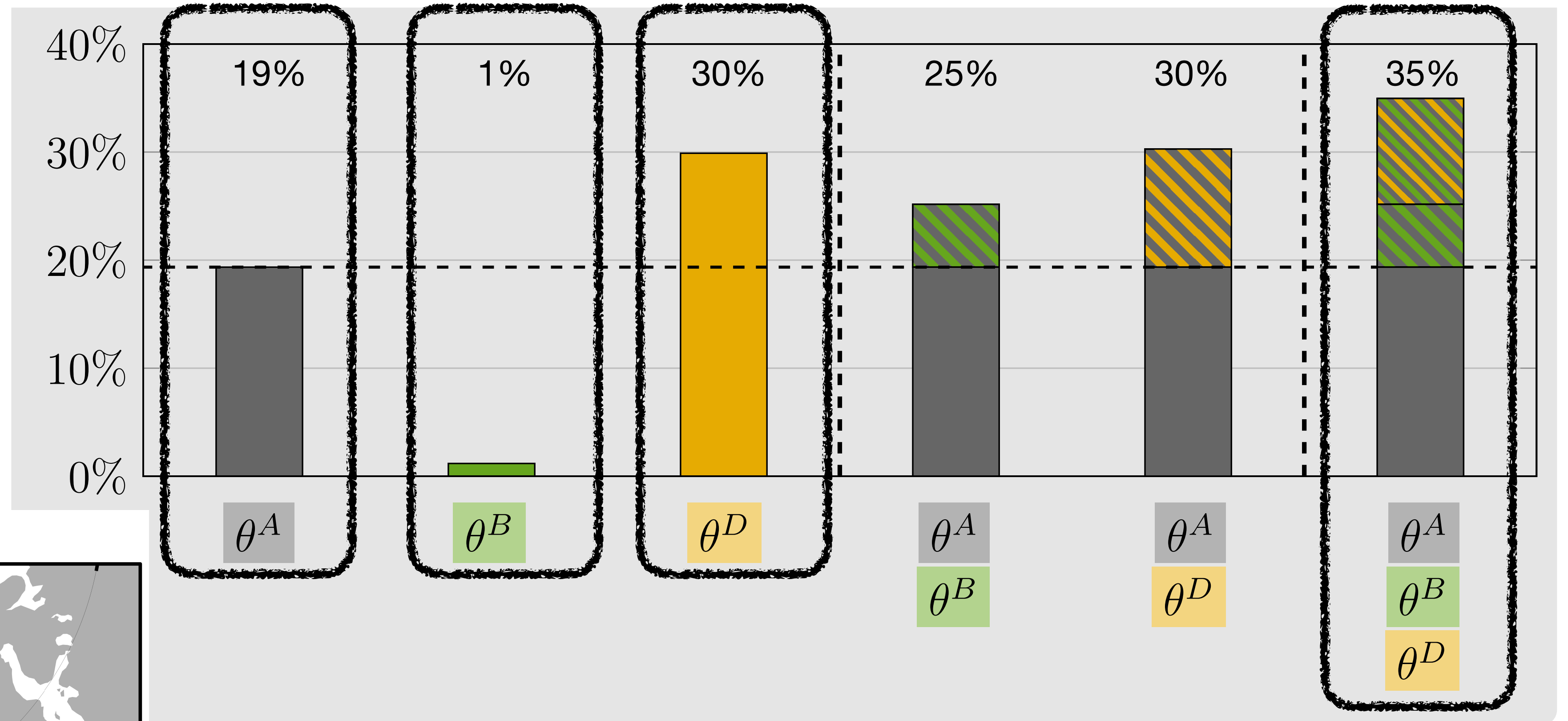
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- Observation A (in the Irminger Sea) and observation B (off Portugal) show **data complementarity**

Assessing data redundancy

Uncertainty reduction in HT_{ISR} by noise-free observations

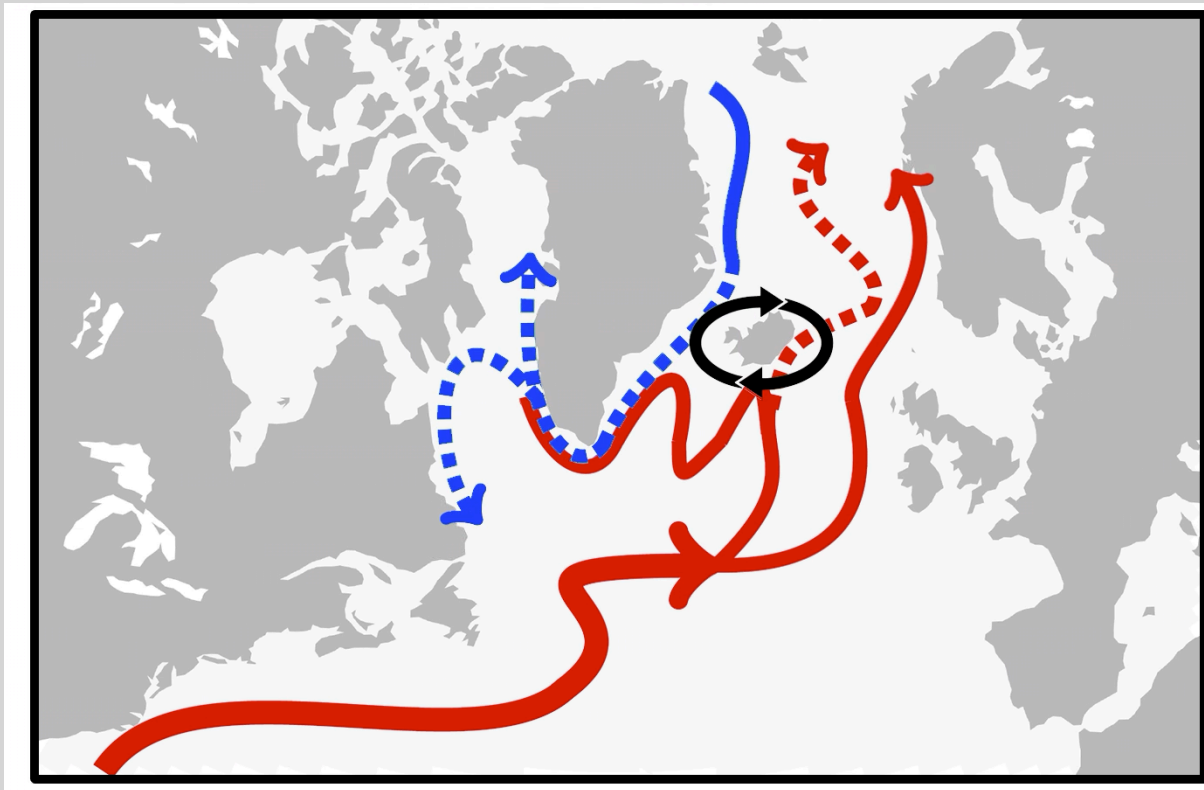


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Summary

Loose et al., *JGR Oceans* (2020)

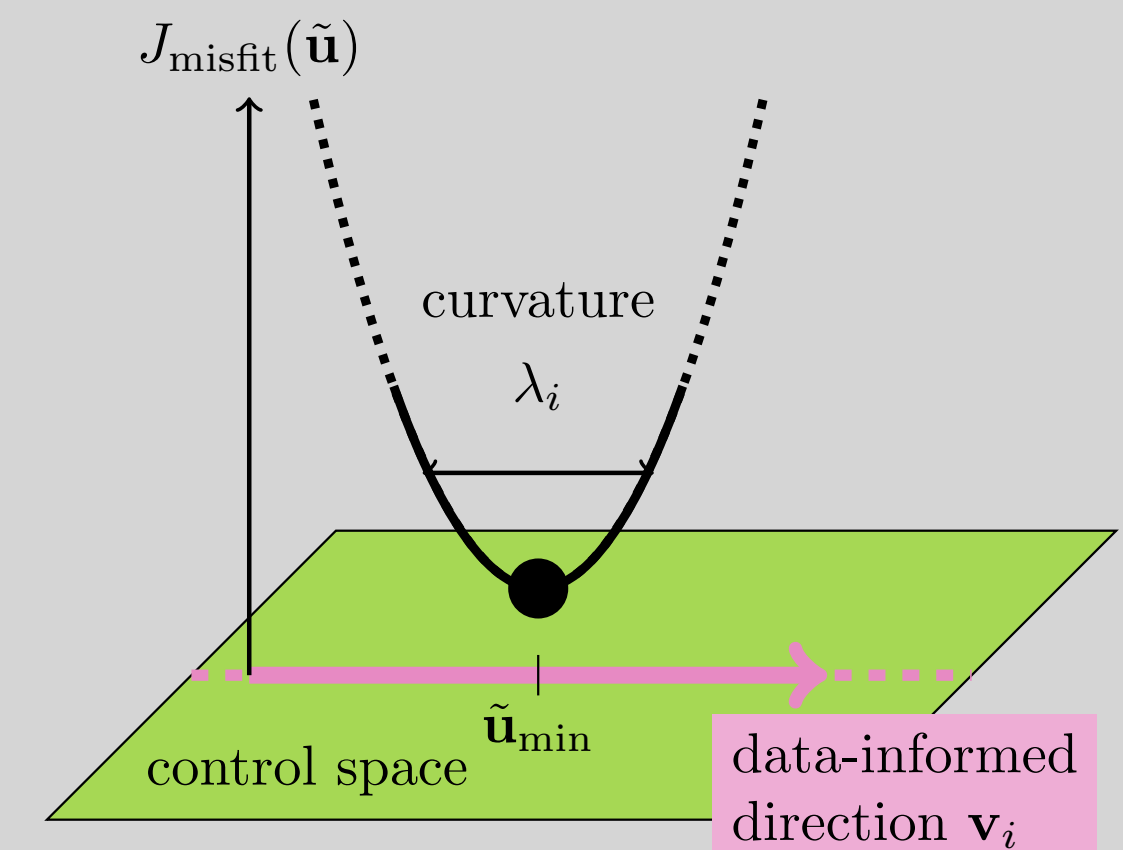
Oceanic teleconnections



Shared adjustment mechanisms

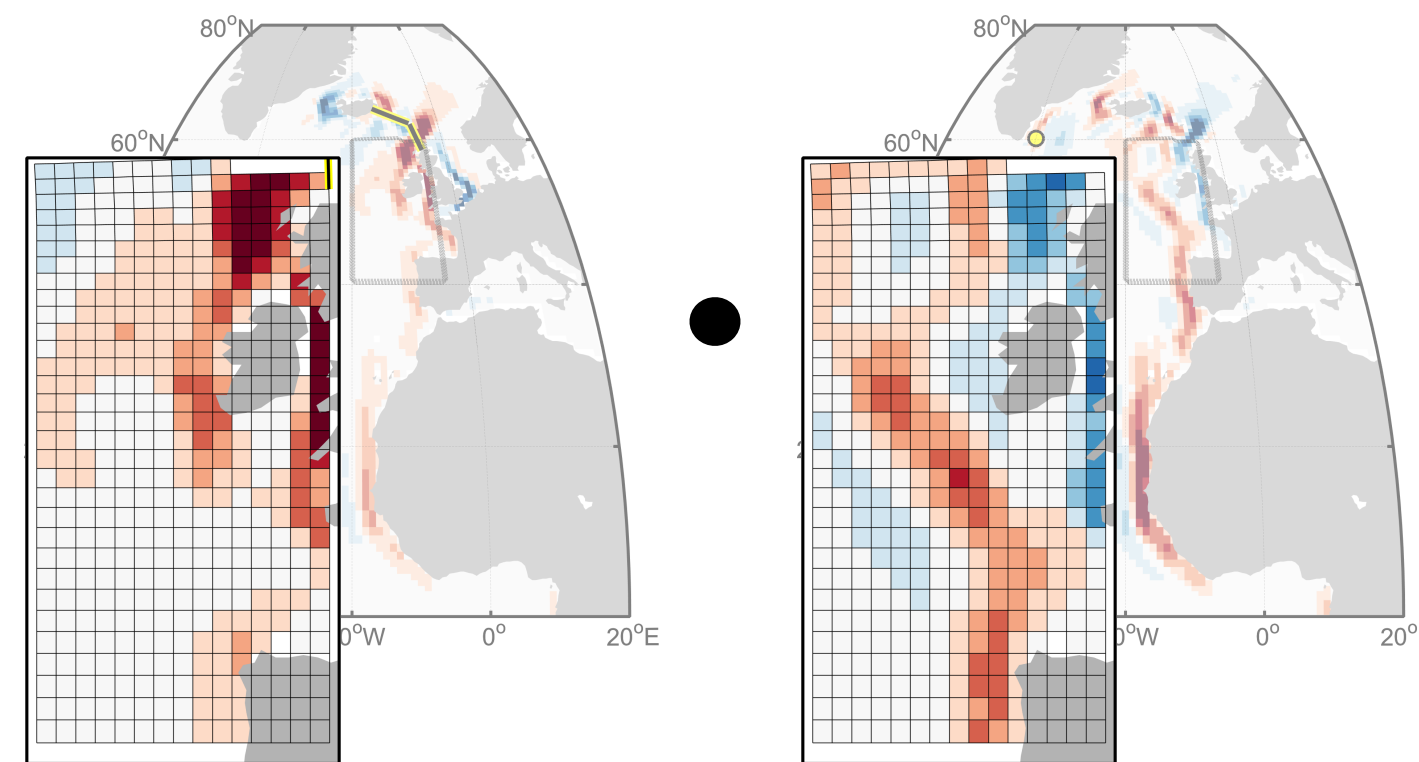
Loose & Heimbach, *JAMES* (2021)

Uncertainty Quantification (UQ)



Data redundancy & complementarity

Adjoint model

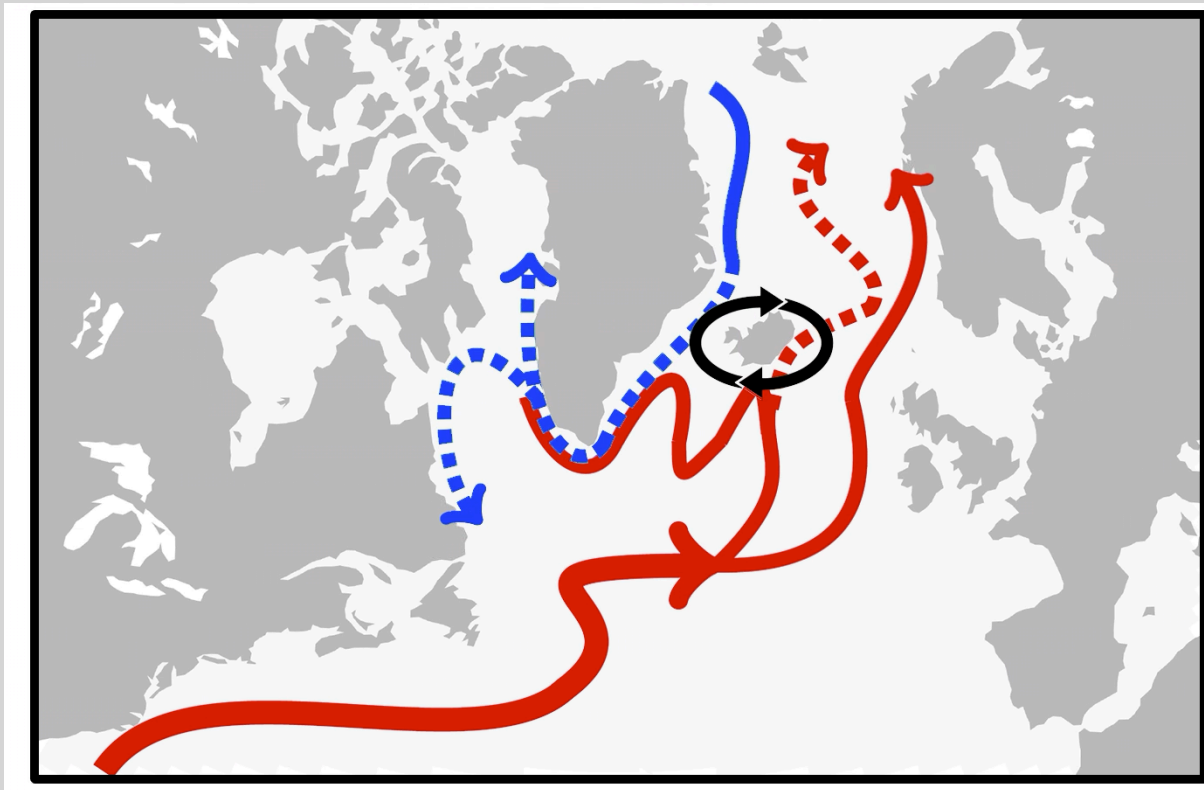


Dynamics-informed & quantitative observing system design

Summary

Loose et al., *JGR Oceans* (2020)

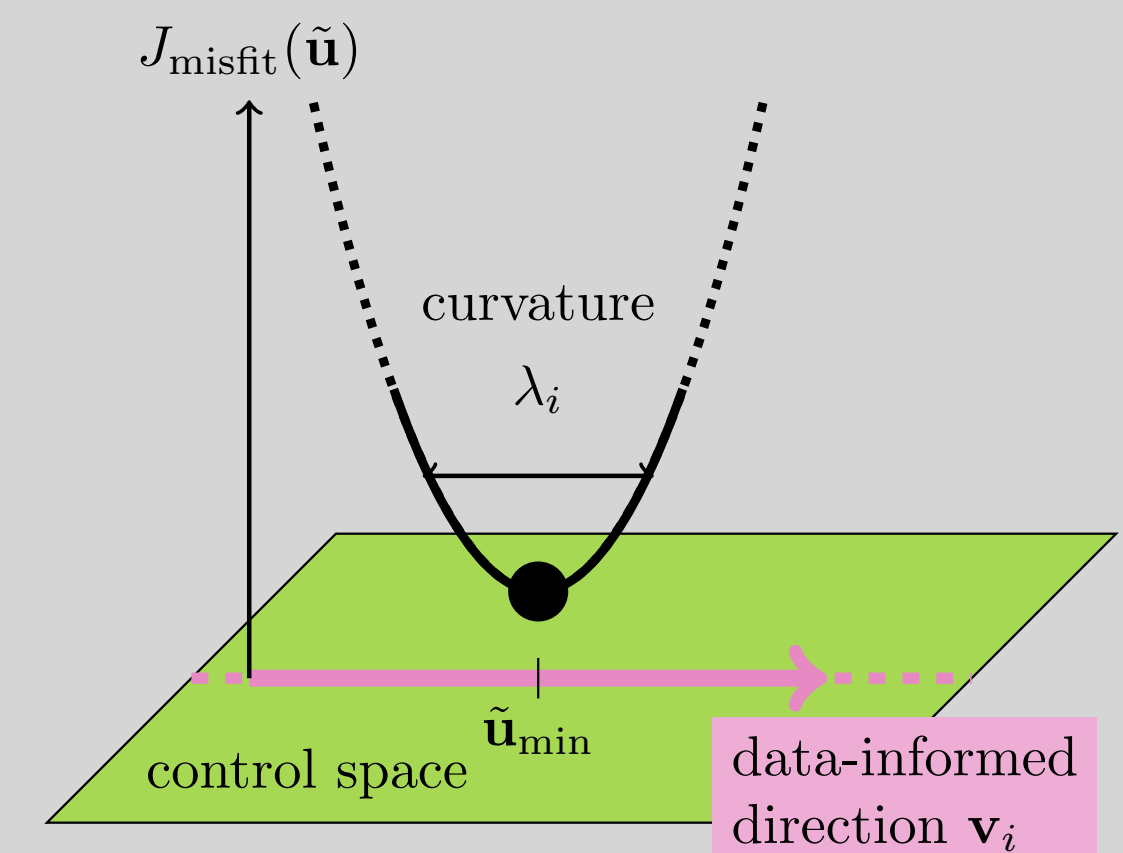
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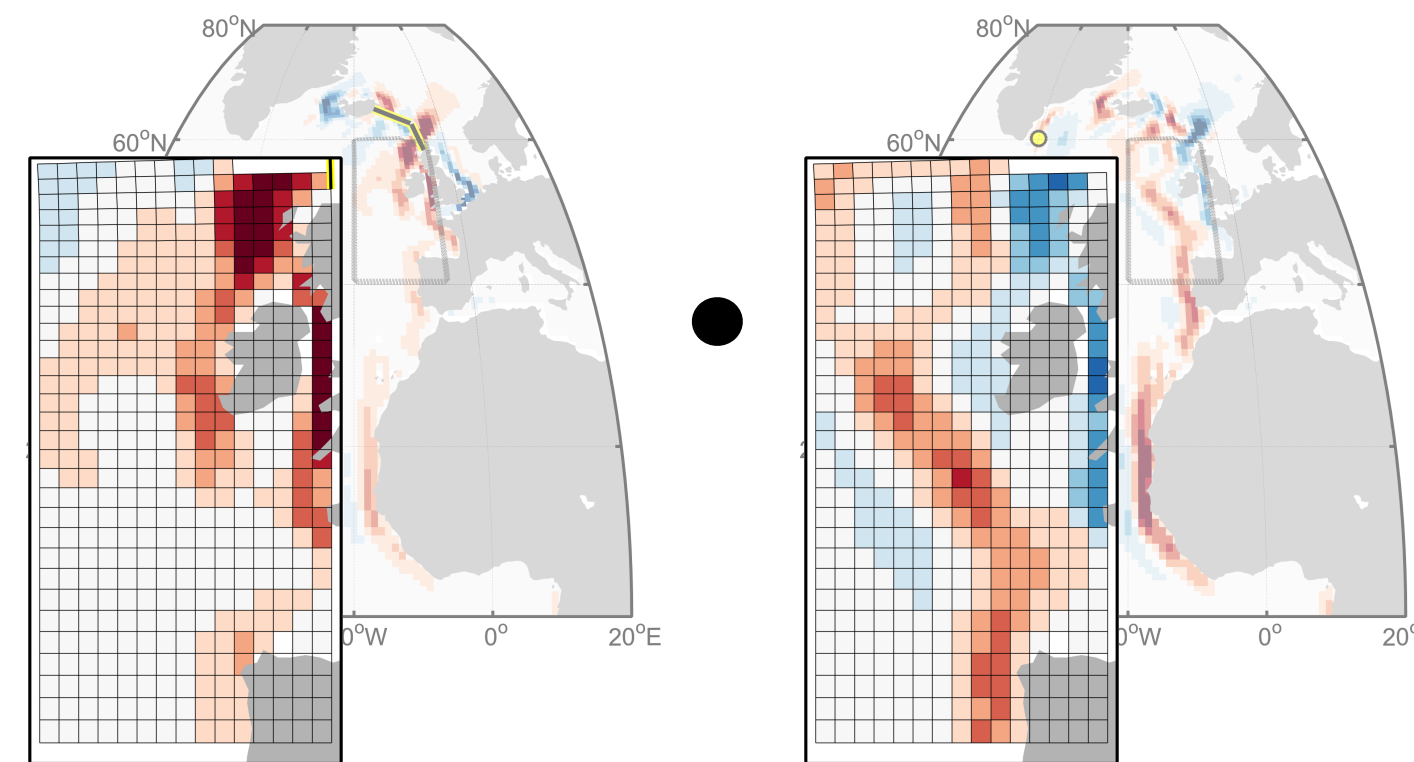
Loose & Heimbach, *JAMES* (2021)

Uncertainty Quantification (UQ)



Data redundancy & complementarity

Adjoint model



Dynamics-informed & quantitative observing system design

Limitation: Adjoint only provides linearized approximation of ocean dynamics
Appropriate for: Large-scale dynamics and Gaussian approximation of uncertainty
Outlook: How to deal with very nonlinear dynamics & implied uncertainty?



Outlook

DJ4Earth :

Differentiable programming in Julia for Earth system modeling

<https://dj4earth.github.io/>

$\mathcal{J}(\mathcal{M}(\bar{u})) = \mathcal{J}(\mathcal{M}_\lambda(\mathcal{M}_{\lambda-1}(\dots(\mathcal{M}_\lambda(\dots(\mathcal{M}_1(\mathcal{M}_0(\bar{u})))))))$ (7.8)

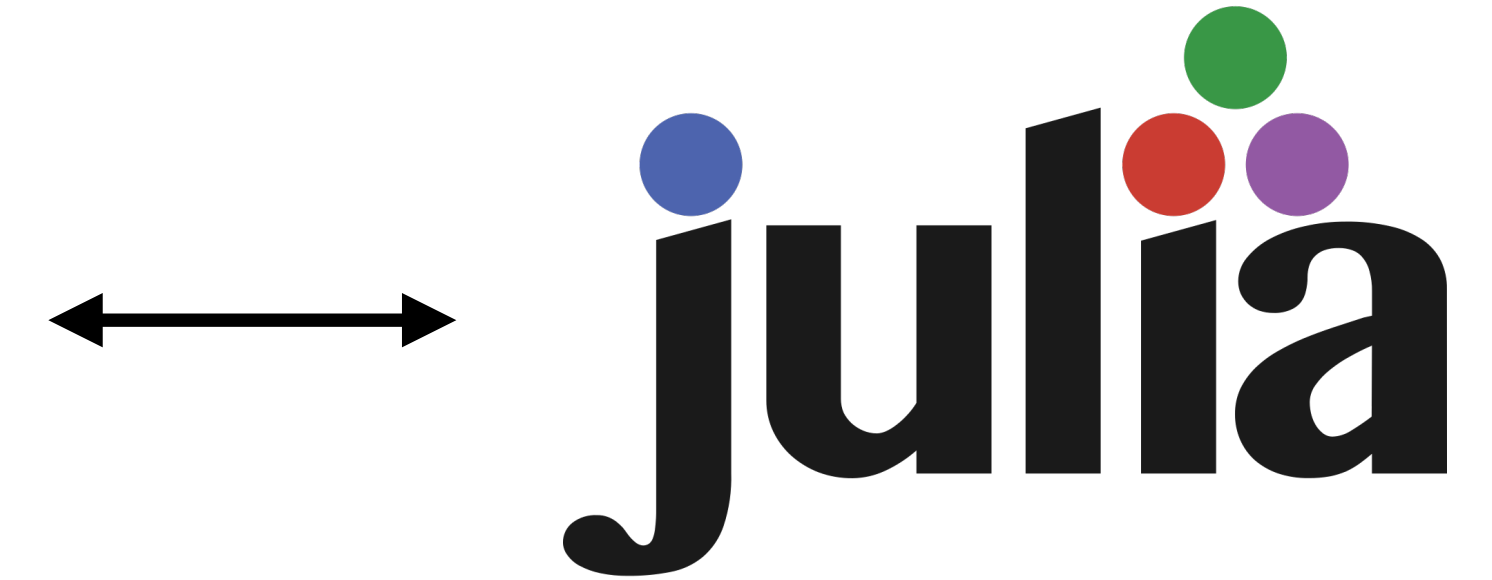
Then, according to the chain rule, the forward calculation reads, in terms of the Jacobi matrices (we've omitted the $\bar{\cdot}$'s which, nevertheless are important to the aspect of tangent linearity; note also that by definition $\langle \nabla_u \mathcal{J}^T, \delta \bar{u} \rangle = \nabla_u \mathcal{J} \cdot \delta \bar{u}$)

$\nabla_u \mathcal{J}(\mathcal{M}(\delta \bar{u})) = \nabla_u \mathcal{J} \cdot \mathcal{M}_\lambda \cdot \dots \cdot \mathcal{M}_\lambda \cdot \dots \cdot \mathcal{M}_1 \cdot \mathcal{M}_0 \cdot \delta \bar{u}$ (7.9)
 $= \nabla_u \mathcal{J} \cdot \delta \bar{u}$

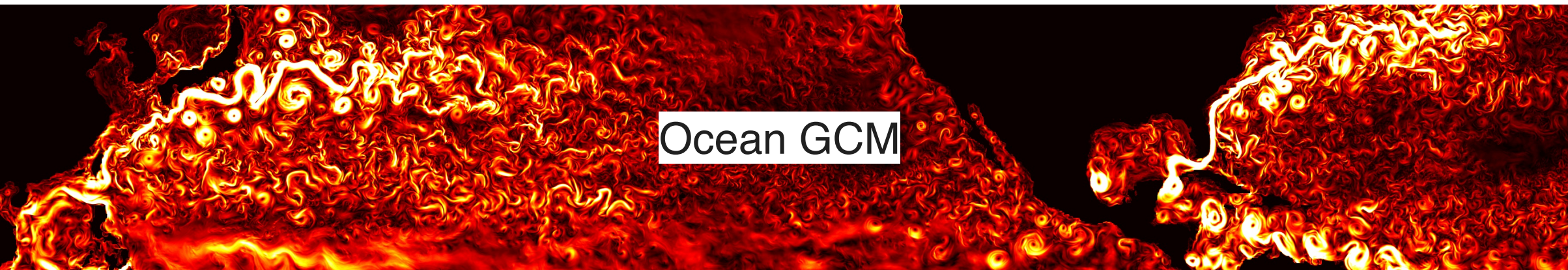
whereas in reverse mode we have

$M^T(\nabla_u \mathcal{J}^T) = M_0^T \cdot M_1^T \cdot \dots \cdot M_\lambda^T \cdot \dots \cdot M_\lambda^T \cdot \nabla_u \mathcal{J}^T$ (7.10)
 $= M_0^T \cdot M_1^T \cdot \dots \cdot \nabla_{u_0} \mathcal{J}^T$
 $= \nabla_u \mathcal{J}^T$

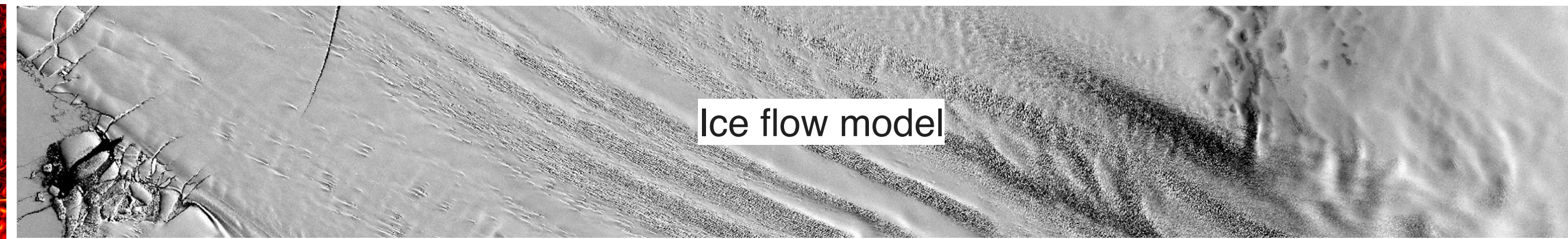
$= \Delta^u \mathcal{J}_\lambda$
 $= \mathcal{V}_\lambda^0 \cdot \mathcal{V}_\lambda^1 \cdot \dots \cdot \Delta^u \mathcal{J}_\lambda$
 $\mathcal{V}_\lambda(\Delta^u \mathcal{J}_\lambda) = \mathcal{V}_\lambda^0 \cdot \mathcal{V}_\lambda^1 \cdot \dots \cdot \mathcal{V}_\lambda^{\lambda-1} \cdot \Delta^u \mathcal{J}_\lambda$ (7.10)



Earth system flagship applications



Ocean GCM



Ice flow model



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Then, according to the chain rule, the forward calculation reads, in terms of the Jacobi matrices (we've omitted the $\bar{\cdot}$'s which, nevertheless are important to the aspect of tangent linearity; note also that by definition $\langle \nabla_u \mathcal{J}^T, \delta \bar{u} \rangle = \nabla_u \mathcal{J} \cdot \delta \bar{u}$)

$\nabla_u \mathcal{J}(\mathcal{M}(\delta \bar{u})) = \nabla_u \mathcal{J} \cdot \mathcal{M}_\lambda \cdot \dots \cdot \mathcal{M}_\lambda \cdot \dots \cdot \mathcal{M}_1 \cdot \mathcal{M}_0 \cdot \delta \bar{u}$ (7.9)

$= \nabla_u \mathcal{J} \cdot \delta \bar{u}$

whereas in reverse mode we have

$M^T(\nabla_u \mathcal{J}^T) = M_0^T \cdot M_1^T \cdot \dots \cdot M_\lambda^T \cdot \dots \cdot M_\lambda^T \cdot \nabla_u \mathcal{J}^T$ (7.10)

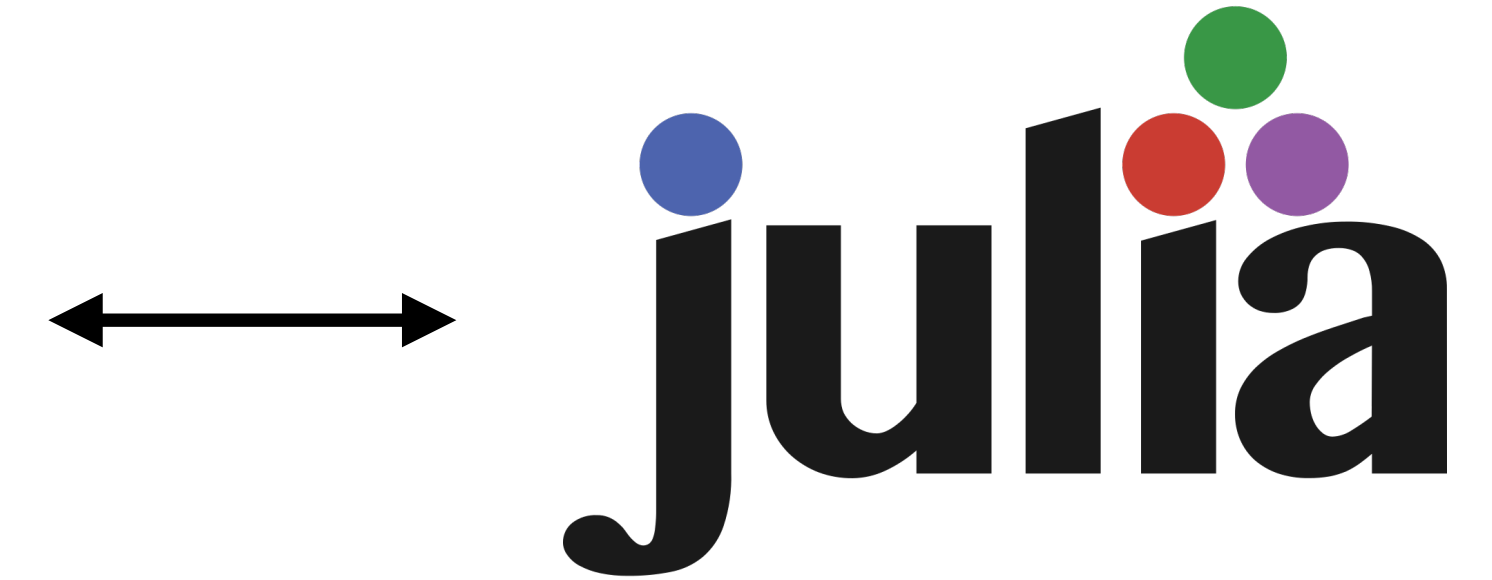
$= M_0^T \cdot M_1^T \cdot \dots \cdot \nabla_u \mathcal{J}^T$

$= \nabla_u \mathcal{J}^T$

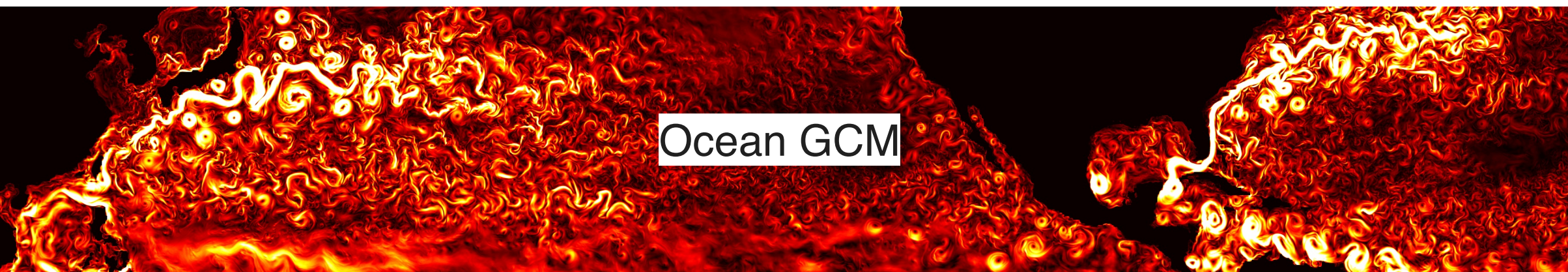
$= \Delta^{\mu} \Delta_{\lambda}$

$= \mathcal{N}_0^{\mu} \cdot \mathcal{N}_1^{\mu} \cdot \dots \cdot \Delta^{\mu} \Delta_{\lambda}$

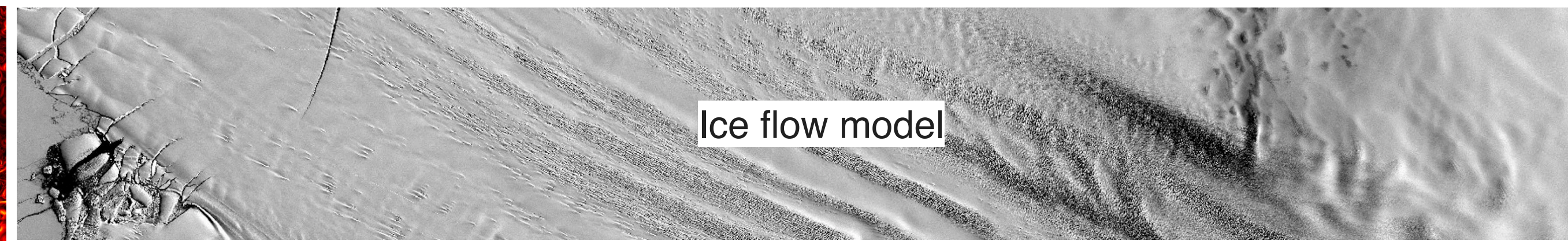
$\mathcal{N}_\lambda(\Delta^{\mu} \Delta_{\lambda}) = \mathcal{N}_\lambda^{\mu} \cdot \mathcal{N}_\lambda^{\mu} \cdot \dots \cdot \mathcal{N}_1^{\mu} \cdot \Delta^{\mu} \Delta_{\lambda}$ (7.10)



Earth system flagship applications



Ocean GCM



Ice flow model

Goal: Advance automatic differentiation (AD) to generate adjoint and back-propagation operators in Earth system models

Julia:

Physics model (PDE) & its adjoint

Earth system model

Seamless integration

Machine Learning (ML) & differentiable programming



Outlook

DJ4Earth :

Differentiable programming in Julia for Earth system modeling

<https://dj4earth.github.io/>

$\mathcal{J}(\mathcal{M}(\bar{u})) = \mathcal{J}(\mathcal{M}_\lambda(\mathcal{M}_{\lambda-1}(\dots(\mathcal{M}_\lambda(\dots(\mathcal{M}_1(\mathcal{M}_0(\bar{u})))))))$ (7.8)

Then, according to the chain rule, the forward calculation reads, in terms of the Jacobi matrices (we've omitted the $\bar{\cdot}$'s which, nevertheless are important to the aspect of tangent linearity; note also that by definition $\langle \nabla_v \mathcal{J}^T, \delta \bar{v} \rangle = \nabla_v \mathcal{J} \cdot \delta \bar{v}$)

$\nabla_v \mathcal{J}(\mathcal{M}(\delta \bar{u})) = \nabla_v \mathcal{J} \cdot \mathcal{M}_\lambda \cdot \dots \cdot \mathcal{M}_\lambda \cdot \dots \cdot \mathcal{M}_1 \cdot \mathcal{M}_0 \cdot \delta \bar{u}$ (7.9)

$= \nabla_v \mathcal{J} \cdot \delta \bar{u}$

whereas in reverse mode we have

$M^T(\nabla_v \mathcal{J}^T) = M_0^T \cdot M_1^T \cdot \dots \cdot M_\lambda^T \cdot \dots \cdot M_\lambda^T \cdot \nabla_v \mathcal{J}^T$ (7.10)

$= M_0^T \cdot M_1^T \cdot \dots \cdot \nabla_{v_0} \mathcal{J}^T$

$= \nabla_v \mathcal{J}^T$

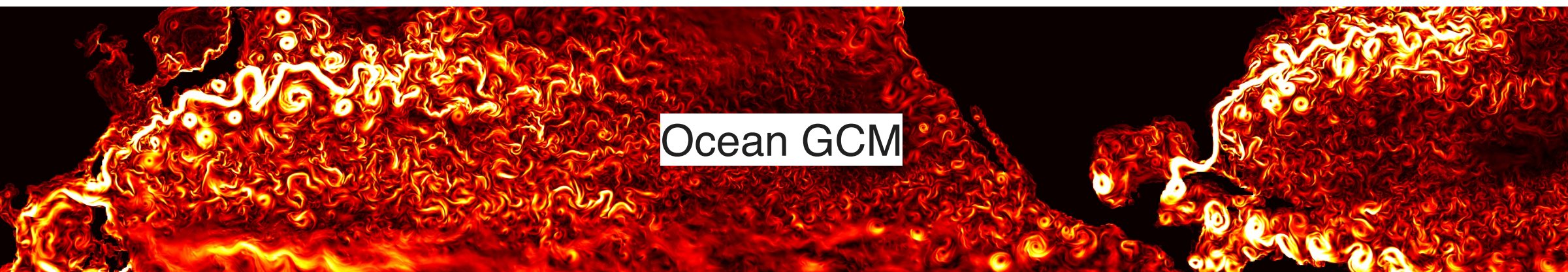
$= \Delta^v \Delta_{\lambda_1}$

$= \mathcal{N}_0^v \cdot \mathcal{N}_1^v \cdot \dots \cdot \Delta^{v_0} \Delta_{\lambda_1}$

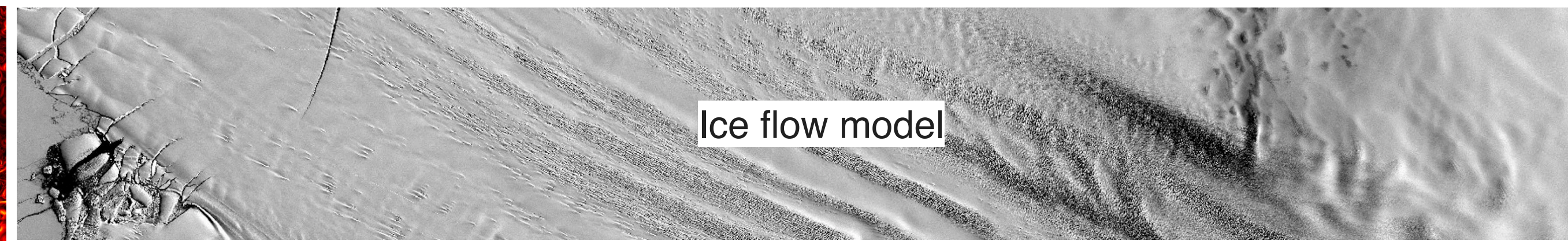
$\mathcal{N}_\lambda(\Delta^v \Delta_{\lambda_1}) = \mathcal{N}_\lambda^v \cdot \mathcal{N}_\lambda^v \cdot \dots \cdot \mathcal{N}_1^v \cdot \Delta^v \Delta_{\lambda_1}$ (7.10)



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Hybrid approaches:

- ML-accelerated sampling for non-Gaussian UQ
- Derivative- and physics-informed ML